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ASSESSING THE IMPACTS OF CLIMATE CHANGE ON THE SURFACE TEMPERATURE OF INLAND LAKES IN MICHIGAN

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ASSESSING THE IMPACTS OF CLIMATE CHANGE ON THE SURFACE TEMPERATURE OF INLAND LAKES IN MICHIGAN

By

Kaitlin L. Reinl

A THESIS

Submitted in partial fulfillment of the requirements for the degree of

MASTER OF SCIENCE

In Applied Ecology

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2016

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This thesis has been approved in partial fulfillment of the requirements for the Degree of MASTER OF SCIENCE in Applied Ecology.

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To my goddaughters Emma June and Rosemary Louise.

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Preface

This thesis is original and unpublished work based on intellectual concepts conceived by Dr. Noel Urban. The results of this thesis will be submitted to peer-reviewed journals for publication. All work, data analyses, and manuscript composition were conducted by Kaitlin Reinl. Dr. Joseph Wagenbrenner, Dr. Noel Urban, and Michael Hyslop provided oversight and assistance in data analysis, experimental design, and manuscript composition.

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The laughs, friendship, and support provided by my best friends, Gretchen Thomas and Kristen McNamara, helped me to get through tough times and made the good times even better. I will always cherish you both.

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Abstract

The aim of this study was to validate and apply a lake model for predicting the susceptibility of small inland lakes in Michigan to changes in thermal regime and increased cyanobacteria growth as a result of future climate conditions. The Freshwater Lake Model was selected, tested for sensitivity to various inputs, and validated through comparison to observed conditions. The sensitivity analysis showed that the lake model was most sensitive to solar radiation, air temperature, and air humidity. Comparison of predicted climate data with observed conditions revealed highly variable climate model error. The lake model validation was conducted using 10 lakes in Wisconsin with observed and modeled meteorological data from 1998 through 1999. The model was valid for predicting surface water temperature, but not for mean temperature, and modeling proceeded with only surface water temperature. The lake model validation resulted in over-prediction when using modeled climate data inputs, which is likely due to inaccuracy in the climate model.

The study area included 517 inland lakes in Michigan. These lakes were divided into 27 groups based on climate, size, and trophic state. Thirteen lake groups were modeled on a daily time step from 2020 to 2099 using prototype lakes and regionally downscaled, modeled climate data. The climate parameters forcing the lake model predictions were analyzed for long-term trends and differences across climates, lake size, and trophic state. The trends in surface water temperature for the entire period and each season from 2020 to 2099 were significant for all modeled lake groups, and lake model surface temperature predictions closely followed modeled air temperature. For all lake groups, the largest increases in surface temperature were observed in spring while the smallest increases occurred in winter. No statistical differences in long-term trends of surface temperature were found between any of the groups regardless of location, size, or trophic state. We analyzed the relationship between changes in periods of minimum and optimum algal growth conditions and climate, lake size, and trophic state. The largest increase in the period with surface temperature above minimum growth temperatures was predicted for small, oligotrophic lakes in the southern Lower Peninsula. This result can mainly be attributed to inherently warmer temperature earlier in the year in more southern latitude positions and the quicker response of small lakes to warming temperatures in comparison to larger lakes. The largest increase in the period with surface temperature above optimal growth temperatures was predicted for large, oligotrophic lakes in the Upper Peninsula. The predicted increase in the number of days the surface temperatures exceeded the optimum growing temperature in the colder Upper Peninsula was greater because of the relatively low number of days at the onset of the modeling period in comparison to lakes in more southern latitudes, and large lakes are able to uptake more heat for longer periods of time. The results of this study illustrated the future trends in surface water temperature and the potential implications for cyanobacteria growth, and can be used to develop plans to prevent and mitigate the spread of cyanobacteria as a result of climate change.

1. Introduction

1.1 Motivation

Small lakes are important in our daily lives. They provide a wealth of ecosystem services such as improving water quality, transporting nutrients, supporting biodiversity, and providing drinking water. In doing so, they support recreation, economic growth, and public health (1). Small inland lakes comprise more than twice the land surface area of large lakes and are much less represented in scientific literature than large lakes (2,3).

Changes in lake temperature due to climate change may greatly alter lake ecosystem composition and function (4,5,6). Lake warming results in decreasing plankton body size and a shift in species towards smaller organisms (7,8). Phytoplankton growth is highly dependent on temperature, and future changes in temperature may affect the competition of phytoplankton communities and result in a loss of diversity in freshwater lakes (9,10,11,12). A study including 143 lakes ranging from northern Europe to South America showed that there is an increase in the percentage of phytoplankton biomass that is cyanobacteria as climate becomes warmer (13). Cyanobacteria can grow at a minimum of 15°C (14,15). Optimum conditions for algal growth occur at 25°C or greater, and as temperature increases the growth rates of blue-green algae exceed those of green algae (14,15). Warmer lake water temperatures may result in a longer duration of optimum conditions for cyanobacteria, resulting in higher potential for outbreaks of harmful algal blooms, even where physical factors have not supported algal blooms in the past (12, 16). Common varieties of bloom-forming blue-green algae that pose a threat include the genera Anabaena, Aphanizomenon, Microcystis, and Oscillatoria (17,14). Each of these blue green algae produces liver toxins, and all except *Microcystis* produce neurotoxins (17). While these algae have capabilities of producing toxins, there are a variety of factors that regulate the amount and potency of the toxic byproducts produced, such as chlorophyll levels, lake mixing, and the ratio of toxin to non-toxin producing bacteria in a given bloom (17).

1.2 Lake Thermal Structure

The thermal structure of lakes is controlled by lake properties and external atmospheric forcing (18, 19). The majority of heat inflows and outflows in a lake occur at the water surface (18, 19). Heat exchanged with sediments can play an important role in very shallow lakes, but is negligible for most water bodies (18, 19). The main components of the lake surface heat exchange are net radiation, latent heat flux, sensible heat flux, sediment heat exchange, and heat storage in the lake (Figure 1.1) (18, 19).

Net radiation includes incoming solar radiation, incoming longwave radiation from the atmosphere, and longwave radiation emitted by the lake surface (18, 19). Incoming solar radiation absorbed by the lake and the amount of incoming radiation is affected by seasonality, latitude position, angle of the sun, scattering of sunlight in the atmosphere by dust and clouds, and the albedo of the lake surface (18, 19). The albedo of a lake is affected by ice and snow cover as well as water clarity (20, 21). Snow and ice cover increase the amount of reflected radiation and increasing water clarity decreases reflectance, generally speaking (20, 21). The effect of water clarity is complicated by the effect of spectral properties of phytoplankton communities because the reflectance is affected by which types of phytoplankton are present in the lake (21). For example, blue algae tend to absorb more light than green algae (21).

Latent heat is exchanged through evaporation and condensation, and sensible heat is exchanged through conduction and convection (18, 19). Latent heat is affected by the air temperature, surface water temperature, vapor pressure gradients, wind speed, and the presence of ice and snow (19). Sensible heat is affected by air and lake temperature as well as wind speed and the presence of ice and snow (18, 19). The relative importance of non-radiation terms in the heat budget for a given lake depends on seasonal variation and lake properties such as geographic location, lake size, and morphology (18). Lake heat storage and distribution in the water column is also affected by these factors (18) and determines the amount of heat available for exchange to the atmosphere. Heat uptake by lakes has also been found to be increase with surface area, mean depth, and basin slope (18, 22).

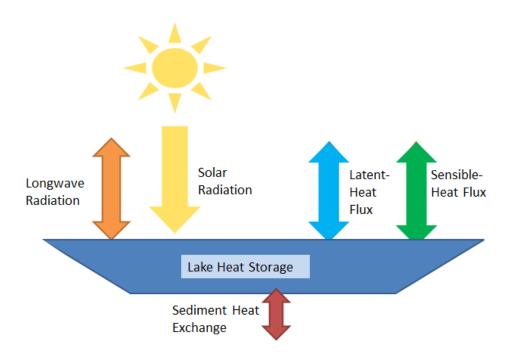


Figure 1.1. Major components of the lake heat budget (Adapted from 23).

The exchange of heat at the lake surface affects temperatures in the lake which affects the density of water (18). Water is most dense at approximately 4°C and becomes less dense as water is warmed or cooled (18,20). This property plays a key role in lake stratification and mixing (18). In dimictic lakes, or lakes that mix twice per year, the lake becomes completely mixed during spring and fall, and the temperature of the water column is uniform at approximately 4°C (19,20). During the summer and winter the lake is stratified into the upper layer, the epilimnion, and the bottom layer, the hypolimnion (19,20). These layers are divided by an area of rapid water temperature change with depth, called the thermocline (19,20). The difference between surface and bottom temperature is larger in summer than in winter. From spring into summer, the surface of

the lake begins to warm from increased solar radiation and air temperatures. The lake begins to stratify, and over the course of the summer the well mixed epilimnion depth increases and the denser, cooler water sinks to the bottom of the lake (19). From summer to fall, the air temperature cools and solar radiation inputs begin to decrease resulting in a net heat loss from the lake (19). Large wind storms aid in breaking down of stratification in late summer (20), and the water column cools to 4°C and again becomes well-mixed during the fall (19,20). Finally, from fall to winter the lake surface cools and the lake becomes stratified again and the cooler layer sits on top of the warmer, denser layer (19, 20). The strength of stratification in winter depends on how cold surface temperature become (19). If the surface cools to 0°C, ice may form (18) and strengthen stratification in the lake (19).

1.3 Previous Work

In a study evaluating contemporary evidence of the effects of climate change on lakes, temperature data were collected from 1990 to 2012 for 142 lakes in Wisconsin (24). An average warming trend of 0.042° C per year was observed across all lakes. Temperature increases in larger lakes (>0.5 km²) were found to be uniform in the water column while smaller lakes (<0.5 km²) were more affected at shallow depths. Larger lakes exhibited a higher median rate of increase in temperature than small lakes for all depths.

In a lake modeling study, the effect of future climate change on four Wisconsin lakes was evaluated using the Dynamic Reservoir Simulation Model (DYERSM) (25,26). DYRESM was run on a daily time step for the years 1986, 1987, and 1989 for three northern lakes and 1921, 1971, and 1972 for one southern lake using current and doubled CO₂ climate scenarios. The model predicted an average increase in surface water temperature of approximately 3°C and average increases in mean temperature ranging from 1°C to 7°C. An increase in ice-free period was observed in northern lakes as well as increased stability in stratification in summer months.

The results of a lake modeling study in Minnesota from 1955 to 1979 using the lake model MINLAKE (27) under doubled CO₂ climate conditions predicted an average increase of 3° C in lake epilimnion temperature with the greatest increase taking place in the fall and spring and the least in summer and winter (28). Maximum surface water temperature was predicted to be higher in southern lakes than northern lakes. Mean water temperature was found to be affected more in oligotrophic lakes than eutrophic lakes due to higher light penetration.

Fang and Stefan (29) simulated the effect of climate change on the water temperature of 27 lakes type in the United States. The Minnesota Lake Model (MINLAKE96) (30) was run on a daily time step from 1961 to 1979 under observed and doubled CO₂ climate conditions. The model predicted an average increase of 3.3° C and 2.6° C in surface and bottom temperatures, respectively (29). The study also showed longer periods of stratification with increased differences in surface and bottom temperatures of up to 3.2° C.

Another modeling study was conducted for the period of 1961 to 2100 under the IPCC B2 emissions scenario (31) using the Freshwater Lake Model (FLake) (32, 33, 34). The B2 emissions scenario corresponds to a moderate increase in greenhouse gas emissions (31). The results showed that lake water temperatures closely followed local

air temperature trends in one polymictic and dimictic lake in Germany (35). A mean water temperature increase of 0.3°C per decade was observed for both lakes. However, the trends varied with season with the largest increasing trend observed in winter. The mixing regime of both lakes was predicted to shift to monomictic by the end of the 21st century.

A comprehensive review of previous studies examined data from the Great Lakes Basin and Precambrian Shield regions (36). The review reported that the simulated lake surface water temperature increase ranges from 1°C to 7°C. Studies in the review also predicted that differences between surface and bottom temperatures will increase, creating a steeper thermocline and longer, more stable periods of stratification. A decrease in ice cover was predicted with warmer climates, particularly in southern regions. Simulations of future hypolimnetic temperatures in deep lakes projected changes between a decrease of 6°C and an increase of 8°C.

1.4 Lake Models

Lake models are invaluable tools in assessing the potential impacts of climate change on lakes. Here we will focus on one-dimensional models that represent the vertical temperature structure in a lake. Two main types of one-dimensional lake models are finite-difference (37,38,39,40) and bulk (37,41) models. Finite-difference models capture lake processes with more complexity, but are computationally intensive and require more data inputs (37). Finite-difference models do not assume that the lake is well-mixed (37,38,39,40). The lake is gridded and the heat budget is solved for each grid and then averaged in the horizontal plane for each depth to determine the vertical profile of the lake (37,39,40). This method can be applied for any physical lake parameter of interest. Examples of finite difference models include MINLAKE96 and DYRESM. Bulk models simplify calculations by assuming the lake has two layers in which the upper layer is well mixed (37,41). Bulk models increase computational efficiency at the cost of losing complexity in lake processes and resolution in the predicted thermal profile (37). Bulk lake models are often used for parameterization in numerical weather prediction, where surface temperature is a main concern, and thus, most validation efforts have been aimed at determining the accuracy of surface temperature predictions (37).

One-dimensional lake models including MINLAKE96 (30), Hostetler model (42), LAKE model (43), and FLake were screened for accuracy, accessibility, type and resolution of input data, and outputs. These models were part of the Lake Model Intercomparison Project (LakeMIP) aimed, in part, to compare the performance and range of applicability of one-dimensional lake models in predicting surface water temperatures (37). All of the lake models satisfactorily reproduced observed surface water temperatures for Sparkling Lake in Wisconsin. The minimum correlation of model predictions with observed temperatures was 0.988, with the exception of MINLAKE96 for which correlation was not reported (37). In another phase of the LakeMIP, the models were applied to a shallow, turbid lake in Germany for the open water season in 2003 (38). The models produced root mean squared errors ranging from 0.80 °C to 1.96°C (38).

We chose the lake model FLake for our study. The accuracy of FLake has been documented in a number of studies that show the model adequately predicts lake surface temperature in a variety of lake types at different locations. FLake was used to determine the impact of climate change on the thermal structure of two lakes in Berlin, Germany

(35,44). One of the lakes is dimictic and the other is polymictic, and in both cases FLake was able to adequately reproduce the temperature and mixing regimes. In another study, FLake performed well when coupled with a regional climate model and used to predict surface temperatures and lake mixing for a small, shallow lake (45). FLake was run on a daily time step for the year 2005 for Sparkling Lake and Trout Bog, located in Vilas County, WI. These are dimictic lakes with moderate urban development and mostly forested watersheds. The model produced a Pearson correlation with observed data of 0.98 and a root mean squared error of 3.2° C (45). The performance of FLake was also tested for two African Great Lakes. The model was able to capture observed water temperatures and mixing within reason (46). The study also found that FLake was sensitive to accuracy of climate data and choice of lake depth and water transparency inputs (46).

FLake was chosen for this study due to its accuracy, accessibility, and ease of application, including availability of input data. FLake is publically available in a windows executable format, where minimal coding skills are required. FLake can produce accurate results with minimal data inputs and no model calibration.

2. Objectives and Hypotheses

The aim of this study was to validate and apply a lake model to identify which lakes are most susceptible to changes in thermal regime as a result of climate change. First, a lake model was selected and validated through comparison of lake model predictions to observed conditions. The validated lake model was then used to predict future water temperature conditions. The model predictions were used to assess potential risk for lakes to cyanobacteria growth.

The specific objectives of the study were to: 1) determine the ability of the lake model to reproduce observed conditions; 2) determine whether the magnitude of change in surface water temperature varied across climates, trophic states, and/or lake size; and 3) apply these results to determine the risk of increased cyanobacteria growth in lakes.

We hypothesized that the greatest magnitude of change in water temperature would be observed in northern, oligotrophic lakes. We also hypothesized that small, eutrophic lakes located in warmer climates would show the lowest magnitude of change in water temperature; however, these lakes would also be most susceptible to negative ecological effects.

3. Methods

The study area included 517 inland lakes in Michigan, which were divided into 27 groups based on climate, size, and trophic state. A lake model was selected, tested for sensitivity, and validated for predicting lake water temperatures using observed and downscaled climate data. A set of thirteen lake groups were selected and one prototype lake was developed for each group. The prototype lakes were modeled on a daily time step from 2020 to 2099 using regionally downscaled modeled climate data. The climate parameters forcing the model predictions and the lake model predictions were analyzed

for long-term trends and differences in trends across climates, lake sizes, and trophic states. The relationships between climate, lake size, and trophic state and changes in the period when surface water temperatures were predicted to be above minimum and optimum growth conditions for cyanobacteria were also analyzed for trends for the years 2020 and 2099.

3.1 Study Area

The study area included inland lakes in Michigan, excluding the Great Lakes. We obtained surface area, mean depth, maximum depth, and Secchi depth for 517 lakes from a publicly available database (47), hereafter referred to as the "MSU dataset" (Table 3.1). Lake data were also collected from the USGS National Hydrography Dataset (NHD) (48) for comparison to the number, spatial coverage, and lake size distribution of the MSU dataset (Figure 3.1and Table 3.2).

Table 3.1. Descri	ptive statistics	for the 517	lakes in the	MSU dataset.

	Minimum	Maximum	Mean	Standard Deviation
Surface Area (km ²)	0.04	81.3	2.8	8.5
Mean Depth (m)	0.8	42.5	4.3	3.6
Maximum Depth (m)	2.1	86.9	14.1	8.5
Secchi Depth (m)	0.5	8.7	3.1	1.3

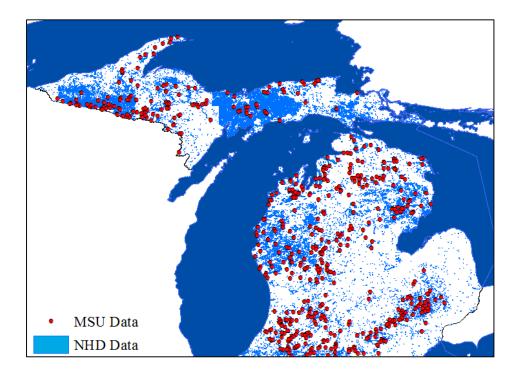


Figure 3.1. Map of lakes from the USGS National Hydrography Dataset (NHD) and data from a Michigan State University study (MSU dataset).

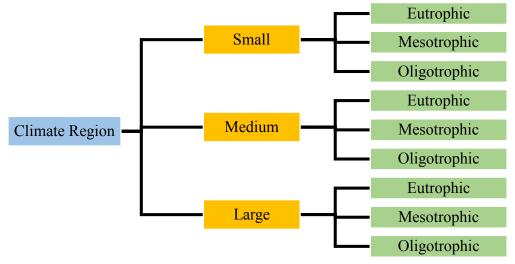
MSU	NHD
517	89995
0.04	0.0003
81.3	86.8
2.8	0.09
8.5	1.1
	517 0.04 81.3 2.8

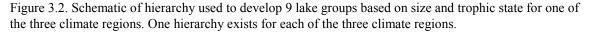
Table 3.2. Summary of descriptive statistics for surface area in MSU and NHD datasets.

3.2 Lake Groups

Lake groups were formed using a hierarchical scheme across climates, surface area, and trophic state, resulting in 27 lake groups for modeling purposes (Figure 3.2). The air temperature data used to determine the climate groups was statistically downscaled for the U.S. (49) and had a spatial resolution of 4.5 km. The long-term mean air temperature was averaged over the period of 1870 to 1999 (50, 51). Climate groups were formed in GIS using natural breaks in air temperature (Figure 3.3), separating the 517 lakes into three climate groups: Upper Peninsula, Northern Lower Peninsula, and Southern Lower Peninsula (Figure 3.3 and Table 3.3). The lakes were also classified by surface area into three groups (small, medium, and large) based on quantiles in GIS (Table 3.3). Secchi depth was used to classify trophic state as either eutrophic, mesotrophic, or oligotrophic (52) for each lake (Table 3.3). The naming convention for the 27 groups is shown in Table 3.4 and the 517 lakes and their corresponding groups are shown in Appendix A.

Spatial distributions of lake size and trophic state are shown in Figure 3.4 and Figure 3.5, respectively. The lake size groups are evenly distributed throughout the three climate regions. Lakes that are eutrophic and mesotrophic are also evenly distributed throughout the climate regions; however, there are more oligotrophic lakes in the more northern climate regions.





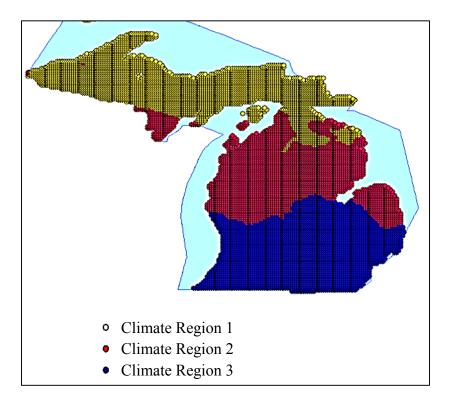


Figure 3.3. Natural breaks in long-term mean temperature in Michigan used to develop climate groups where Climate Region 1, 2, and 3 are the Upper Peninsula, northern Lower Peninsula, and southern Lower Peninsula, respectively.

Table 3.3. Summar	y of criteria used to	develop lake classification scheme.
-------------------	-----------------------	-------------------------------------

Factor	Criteria
Climate Class	Temperature (K)
Upper Peninsula	< 278.43
Northern Lower Peninsula	278.43 - 281.01
Southern Lower Peninsula	>281.01
Size Class	Surface Area (km ²)
Small	<0.05
Medium 0.05-1.25	
Large	1.25+
Trophic State	Secchi Depth (m)
Eutrophic	0-2
Mesotrophic	2-4
Oligotrophic	4+

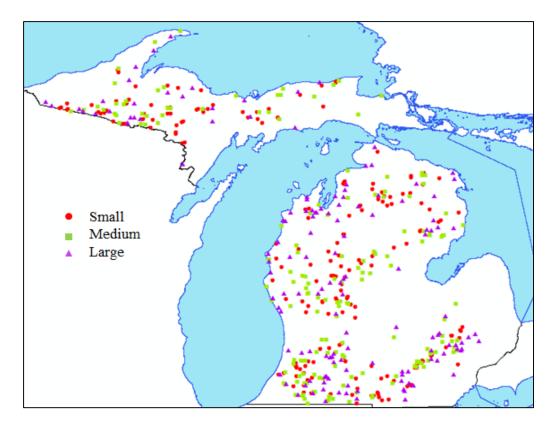


Figure 3.4. Spatial distribution of small, medium, and large lakes in the MSU dataset.

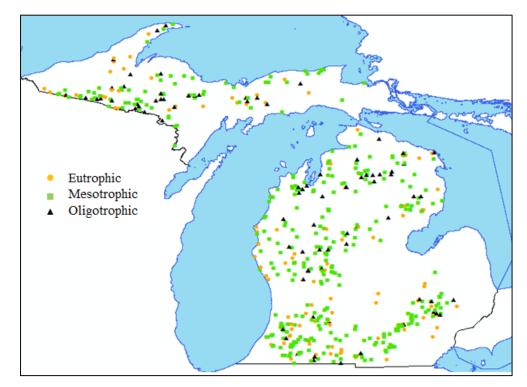


Figure 3.5. Spatial distribution of eutrophic, mesotrophic, and oligotrophic lakes in the MSU dataset.

We grouped the NHD and MSU dataset by climate and lake size to compare lake size distributions (Figure 3.6 and Table 3.3). The NHD did not contain trophic state data and so the datasets could not be compared across trophic states. A larger number of lakes in the NHD are in the Upper Peninsula and more lakes in the MSU dataset are located in the Lower Peninsula. The NHD also has a smaller minimum lake surface area, and larger maximum surface area, than the MSU dataset. Despite the observed differences, the MSU dataset does adequately represent the spatial distribution and size distribution of the larger sample contained in the NHD.

Table 3.4 Naming convention, description, and number of lakes (N) for the 27 lake groups. Lake groups that were modeled and which round they were modeled in are shown in the 'Modeled' column. Superscripts denote round 1 or round 2 modeling.

Group Name	Description	N	Modeled
UP-Sm-E	Upper Peninsula, Small, Eutrophic	14	Y^1
UP-Sm-M	Upper Peninsula, Small, Mesotrophic	39	Ν
UP-Sm-O	Upper Peninsula, Small, Oligotrophic	24	Ν
UP-Md-E	Upper Peninsula, Medium, Eutrophic	8	Y^2
UP-Md-M	Upper Peninsula, Medium, Mesotrophic	29	Ν
UP-Md-O	Upper Peninsula, Medium, Oligotrophic	10	Ν
UP-Lg-E	Upper Peninsula, Large, Eutrophic	10	Y^2
UP-Lg-M	Upper Peninsula, Large Mesotrophic	25	Y^2
UP-Lg-O	Upper Peninsula, Large, Oligotrophic	11	Y^2
NLP-Sm-E	Northern Lower Peninsula, Small, Eutrophic	1	Y^1
NLP-Sm-M	Northern Lower Peninsula, Small, Mesotrophic	41	Ν
NLP-Sm-O	Northern Lower Peninsula, Small, Oligotrophic	8	Ν
NLP-Md-E	Northern Lower Peninsula, Medium, Eutrophic	13	Ν
NLP-Md-M	Northern Lower Peninsula, Medium, Mesotrophic	27	Ν
NLP-Md-O	Northern Lower Peninsula, Medium, Oligotrophic	5	Ν
NLP-Lg-E	Northern Lower Peninsula, Large, Eutrophic	9	Ν
NLP-Lg-M	Northern Lower Peninsula, Large. Mesotrophic	33	Ν
NLP-Lg-O	Northern Lower Peninsula, Large Oligotrophic	8	Y^2
SLP-Sm-E	Southern Lower Peninsula, Small, Eutrophic	13	Y^1
SLP-Sm-M	Southern Lower Peninsula, Small, Mesotrophic	49	Y^1
SLP-Sm-O	Southern Lower Peninsula, Small, Oligotrophic	7	Y^1
SLP-Md-E	Southern Lower Peninsula, Medium, Eutrophic	17	N
SLP-Md-M	Southern Lower Peninsula, Medium, Mesotrophic	49	Y^2
SLP-Md-O	Southern Lower Peninsula, Medium, Oligotrophic	6	Ν
SLP-Lg-E	Southern Lower Peninsula, Large, Eutrophic	18	Y^1
SLP-Lg-M	Southern Lower Peninsula, Large, Mesotrophic	39	Ν
SLP-Lg-O	Southern Lower Peninsula, Large, Oligotrophic	4	Y^2

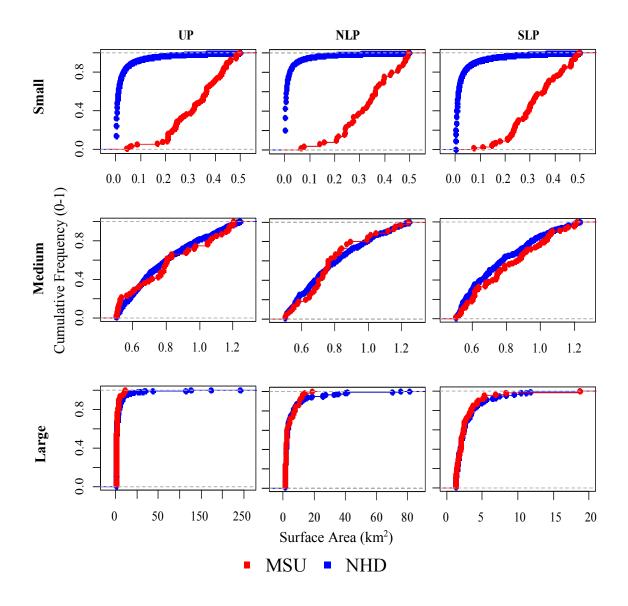


Figure 3.6. Cumulative frequency distributions of lake surface area by size group (rows) and climate groups (columns) for the MSU dataset and NHD. The climate groups are in the Upper Peninsula (UP), northern Lower Peninsula (NLP), and southern Lower Peninsula (SLP)). Note the different surface area scales among the large lake groups.

3.3 Lake Model

FLake is a physically-based, freshwater lake model that predicts lake temperature and mixing as well as ice cover and energy fluxes (32, 33, 34). The main uses of the model to date have been for representing the effect of lakes in numerical weather prediction and as a stand-alone lake model. The model is intended to be used without calibration because it

may reduce the model's predictive capability (33). The model is driven by both meteorological data and lake properties including: solar radiation, air temperature, air humidity, wind speed, cloudiness, latitude, depth, fetch, and extinction coefficient. Secchi depth was used to determine the extinction coefficient using the relationship that the extinction coefficient is 1.7 divided by Secchi depth in meters (53). The model outputs of interest to this study are surface and mean water temperature, but other outputs include sensible and latent heat flux, sediment heat exchange, convective heat flux, and ice and snow cover.

The model divides the lake into two layers in which the upper layer is completely mixed and is confined by the thermocline which is described using the self-similarity concept (54,55). The thermocline shape is determined using a single equation that calculates the shape coefficient as a function of depth and time. As wind speed increases, the thickness of the upper layer deepens and the shape of the thermocline remains the same. The heat budget for the entire water column is governed by the following equation:

$$D\left(\frac{d\overline{\theta}}{dt}\right) = \frac{1}{\rho_{w}c_{w}} [Q_{s} + I_{s} - Q_{b} - I(D)]$$

where $\overline{\theta}$ is mean temperature, t is time, D is the maximum depth, ρ_w is the density of water, c_w is the specific heat of water, Q is the turbulent heat flux, and I is the radiation heat flux. Subscripts of s and b denote the lake surface and bottom, respectively. The traditional components of the lake heat budget that make up Q_s are the latent and sensible heat exchange with the atmosphere. I_s is made up of incoming solar radiation and longwave radiation emitted by the atmosphere, and longwave radiation emitted by the lake surface. Q_b is the heat exchange with the sediments. For a more detailed description of the lake parameterization scheme see Mironov (33).

3.4 Sensitivity Analysis

A sensitivity analysis was conducted for FLake using Allequash Lake (see section 3.5.1 below) and the modeled climate data obtained for the northern validation site (see section 3.5.1 below). Each of the climate model input parameters was varied individually, and the model was run keeping all other variables constant. Climate input variables were increased and decreased by 10% and 50% to determine the effects of different inputs on surface and mean water temperature predictions (Table 3.5). The extinction coefficient was increased and decreased by 50% and the model was run without varying any climate variables. All of the values in Table 3.5 could realistically exist in the natural world; however, combinations of the varied and unvaried parameters may not. Furthermore, each of the climate input parameters were varied by $\pm 10\%$ and $\pm 50\%$, but the amount of natural variability that these changes comprise is different for each climate variable (Table 3.6). For example, the change in the mean solar radiation from a 50% increase makes up 61% of the standard deviation of the original dataset, while for the same increase in air temperature the difference in means only makes up approximately 21% of the standard deviation of the original dataset.

Climate Parameter	Statistic	±0%	+10%	-10%	+50%	-50%
	Minimum	3.1	3.4	2.8	4.7	1.6
Solar Radiation	Maximum	367.9	404.7	331.1	551.9	184.0
(W/m^2)	Mean	114.9	126.4	103.4	172.4	57.5
	Standard deviation	93.8	103.2	84.4	140.7	46.9
	Minimum	-22.7	-25.0	-20.5	-34.1	-11.4
Air Temperature	Maximum	22.8	25.1	20.5	34.2	11.4
(°C)	Mean	4.5	5.0	4.1	6.8	2.3
_	Standard deviation	10.5	11.5	9.4	15.7	5.2
	Minimum	3.3	3.6	3.0	5.0	1.7
Air Humidity	Maximum	46.2	50.8	41.6	69.3	23.1
(mb)	Mean	20.3	22.3	18.3	30.4	10.2
	Standard deviation	10.8	11.9	9.7	16.2	5.4
	Minimum	0.1	0.1	0.1	0.2	0.1
Wind Speed	Maximum	10.7	11.8	9.6	16.1	5.4
(m/s)	Mean	4.0	4.4	3.6	5.9	2.0
	Standard deviation	1.8	2.0	1.6	2.7	0.9
	Minimum	0.0	0.0	0.0	0.0	0.0
Cloudiness	Maximum	1.0	1.0	0.9	1.0	0.5
(0-1)	Mean	0.8	0.9	0.8	1.3	0.4
	Standard deviation	0.2	0.2	0.2	0.3	0.1

Table 3.5. Descriptive statistics of varied climate data for the sensitivity analysis.

Table 3.6. Change of varied climate parameters means relative to the standard deviation in the original dataset for each climate input.

	Change in means relative to standard deviation (%)					
Parameter	10%	-10%	50%	-50%		
Solar Radiation	12	12	61	61		
Air Temperature	5	4	22	21		
Air Humidity	19	19	94	94		
Wind Speed	22	22	100	100		
Cloudiness	50	0	100	100		

3.5 Model Validation

The model's mean and surface water temperature predictions were first validated using observed meteorological conditions from 1998 through 1999. We then conducted a second validation using modeled past climate data from 1998 through 1999 to determine how well the combined climate and lake models predicted lake temperatures. Validation was conducted using regression-based equivalence testing for six lakes located in northern Wisconsin and four lakes in southern Wisconsin.

3.5.1 Validation Lake Data

Lakes were chosen for validation based on location, availability of long-term water temperature and meteorological data, and variability of surface area, depth, and trophic state (Table 3.7). The locations of the northern and southern validation sites are shown in Figure 3.7. Data from 1998 through 1999 for the ten validation lakes were obtained from the North Temperate Lakes-Long-Term Ecological Research Network (56). The water temperature measurements were taken at approximately each meter from the surface to the maximum depth for each lake except Lake Wingra, where measurements extended only to the mean depth. The daily mean temperature of the water column was calculated using weighted averages based on each lake's morphology (57). The morphology for Fish Lake was not available, so the morphology of Crystal Lake was used as a proxy based on similarities in the lake bathymetry. For Lake Wingra, the arithmetic mean was used to calculate the mean water temperature, which may have resulted in an overestimate of the mean temperature.

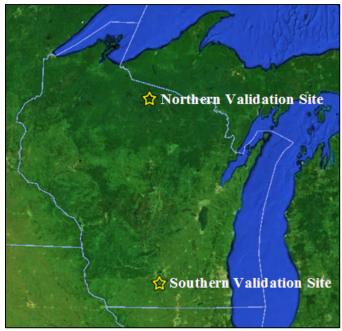


Figure 3.7. Map of the location of the northern and southern validation sites in Wisconsin (Google Earth Map Data: NOAA, Landsat, 58. For image permissions see Appendix D, 59).

Lake	Latitude	Area (km ²)	Max Depth (m)	Extinction Coefficient (1/m)	Fetch (m)	Validation Site
Crystal Bog	46.008	0.01	2.5	1.2	68	Northern
Crystal Lake	46.003	0.38	20.4	0.2	785	Northern
Big Muskellunge Lake	46.021	3.63	21.3	0.3	2380	Northern
Sparkling Lake	46.008	0.64	20.0	0.3	755	Northern
Allequash Lake	46.038	1.64	8.0	0.6	1030	Northern
Trout Lake	46.029	15.65	35.7	0.4	3107	Northern
Fish Lake	43.287	0.80	18.9	0.8	1550	Southern
Lake Mendota	43.099	39.61	25.3	0.6	7503	Southern
Lake Wingra	43.053	1.36	6.7	2.6	1800	Southern
Lake Monona	43.063	13.60	22.5	0.8	5000	Southern

Table 3.7. Summary of lake property data for validation lakes.

3.5.2 Validation Meteorological Data

Observed Meteorological Data

Daily observed meteorological conditions from 1998 through 1999 were obtained for each validation site. Solar radiation, air temperature, air humidity, and wind speed data were obtained from the North Temperate Lakes-Long-term Ecological Research Network (60, 61). Woodruff Airport data were used for lakes at the northern validation site, and data compiled from three observation stations in Madison, WI were used for lakes at the southern validation site (Table 3.7). Fetch was determined by identifying the predominant wind direction at each location and measuring the longest corresponding distance across each lake in GIS.

Modeled Meteorological Data

Downscaled past (1998-1999) and future (2020-2099) meteorological data were obtained from the United States Geological Survey (USGS) Regional Climate Viewer (62). The Regional Climate Viewer produces climate data at 15 km resolution for the eastern United States (62) from a variety of global circulation models. The MPI ECHAM5 climate model (63) with the IPCC A2 emissions scenario (31) was selected to model meteorological conditions for the validation period (1998-1999) as well as for future climate conditions (2020-2099). The A2 emission scenario is one of the highest emission rate scenarios. In this scenario, population continues to increase to over 10 billion by 2050, CO₂ emissions are expected to increase to 870 ppm by the end of the 21^{st} century, methane and nitrous oxide concentrations are expected to increase, and sulfur dioxide increases to a peak in 2050 and then gradually decreases. The MPI ECHAM5 climate model was evaluated with 21 other climate models and was shown to have a lower overall error of more than 10% compared to the typical model error and 30% to 40% less error than the least accurate models in the region where the study site is located (64, 65). The typical model error is the median root square error of all of the models for a given meteorological variable (64).

3.5.3 Validation Statistics

To validate the lake model, we used regression-based equivalence testing. Regression is often used to validate lake models by plotting model predictions and observed data against each other to develop a linear regression. Ideally the slope would equal 1 and the intercept would be equal to 0 if there is complete agreement between model predictions and observed data. Regression-based equivalence testing evaluates the similarity of the regression slope to 1 and intercept to 0. The traditional null and alternative hypotheses are reversed so that the null hypothesis is that the slope and intercept are different from 1 and 0, respectively, and the alternative hypothesis is that the slope and intercept are equal to 1 and 0, respectively (*66*). A regression model is developed for observed and predicted data, and if the slope and intercept fall within the region of similarity, which is based on a-priori values and the significance level, the null hypothesis is rejected.

We used R statistical software (67) and the R package "equivalence" (68) to conduct equivalence testing to validate FLake. The equivalence package uses a bootstrap approach to develop alternative regressions using subsets of the data and determines whether the slope and intercept fall within the predetermined regions of similarity for each regression generated in the bootstrap. To reject the null hypothesis of dissimilarity, the slope and intercept must fall within the region of similarity for at least 95% of the iterations.

The regions of similarity for the lake temperature output using the observed meteorological data were $\pm 15\%$ for the slope and ± 1.0 °C for the intercept. Looser criteria were used to assess the FLake model outputs when driven by modeled meteorological data to accommodate added uncertainty from using modeled meteorological data. The regions of similarity for lake temperature using the modeled climate inputs were $\pm 20\%$ and ± 4.0 °C for the slope and intercept, respectively. The bootstrap test was iterated 1000 times for each lake and a significance level of 0.05 was used for all statistical tests.

3.6 Modeling Approach

We used FLake and modeled future climate data to predict lake temperatures from 2020 to 2099. A set of 13 lake groups were selected for modeling that captured grouping extremes and allowed us to compare lake responses across climates, lake sizes, and trophic states. A prototype lake was developed for each of the selected 13 lake groups and used for modeling future conditions. The midpoint of the latitude range corresponding to each climate group was used for model inputs. For each of the prototype lakes, the model was run on a daily time step for the years 2020 to 2099. The model predictions were analyzed for long-term trends and differences among lake groups. The changes in the period when surface water temperatures were predicted to be above minimum and optimal growth temperatures for cyanobacteria were also analyzed for trends by lake group for the years 2020 and 2099.

3.6.1 Prototype Lake Development and Characteristics

One prototype lake was developed for each lake group by averaging the lake characteristics of all lakes within the group. The fetch of the prototype lake was approximated by the square root of the surface area (69). The prototype lake was used to

model future conditions to represent the response of all the lakes in the group for a given lake type.

We selected a set of seven lake groups from the 27 combinations of climate, size, and trophic state for the first round modeling. One small eutrophic lake from each climate group (UP-Sm-E, NLP-Sm-E, and SLP-Sm-E) was modeled to compare results across climates. Similarly, one small, one medium, and one large eutrophic lake in the warmest climate group (SLP-Sm-E, SLP-Md-E, and SLP-Lg-E) was modeled to compare the effect of lake size. Finally, eutrophic, mesotrophic, and oligotrophic lakes, all small and from the southern-most climate group (SLP-Sm-E, SLP-Sm-E, SLP-Sm-M, and SLP-Sm-O) were modeled to compare trophic states.

Based on a lack of differences among the lakes selected for the first round of modeling, another set of six lakes were modeled to capture more of the extremes of the lake groups. The following lake groups were modeled in the second round: one large, oligotrophic lake in each climate group (UP-Lg-O, NLP-Lg-O, and SLP-Lg-O); one large eutrophic, mesotrophic, and oligotrophic lake in the Upper Peninsula (UP-Lg-E, UP-Lg-M, and UP-Lg-O); and one small, medium, and large eutrophic lake in the Upper Peninsula (UP-Sm-E, UP-Med-E, UP-Lg-E). The characteristics of the thirteen modeled prototype lakes are shown in Table 3.8. Note that some lakes did not have mean depth data which may have resulted in a skewed mean depth for lake groups. An interesting characteristic of the prototype lakes developed from the MSU dataset is that lakes in the small size class tend to have steeper basin slopes than large lakes (Table 3.8).

Lake Group	Ν	Surface Area (km ²)	Mean Depth (m)	Maximum Depth (m)	Secchi Depth (m)	Fetch (m)
UP-Sm-E	14	0.3	3.6 ³	10.0	1.4	576
UP-Md-E	8	0.8	2.7^{1}	6.7	1.1	913
UP-Lg-E	10	5.0	3.1 ²	7.6	1.3	2233
UP-Lg-M	25	5.1	4.0^{2}	11.5	0.6	2266
UP-Lg-O	11	2.4	6.2	17.8	0.3	1547
NLP-Sm-E	1	0.4	NA	7.6	0.9	608
NLP-Lg-O	8	4.1	6.5 ¹	21.0	0.3	2034
SLP-Sm-E	13	0.4	3.54	9.9	0.9	600
SLP-Sm-M	49	0.3	4.7^{10}	11.8	0.6	571
SLP-Sm-O	7	0.3	6.0 ³	16.1	0.4	549
SLP-Md-E	17	0.8	3.9 ²	10.3	1.4	893
SLP-Lg-E	18	3.5	3.7 ²	9.4	1.8	1877
SLP-Lg-O	4	2.7	6.2	28.7	4.5	1642

Table 3.8. Summary of prototype lake characteristics for each modeled lake group. Superscripts in the mean depth column indicate the number of lakes in the set that did not have mean depth data. NA indicates that there were no data available for the group.

3.7 Trend Analysis

The modeled climate data and lake modeling results were analyzed for trends using two procedures. First, the data were analyzed for trend using seasonal decomposition by loess smoothing (70). Seasonal decomposition is used to determine and remove seasonality from a time series dataset to reveal underlying trends. These methods were applied using the "stl" function in the R package "*stats*" (67). The seasonal component was removed and loess smoothing was then applied to the detrended data. Finally, the long-term trend was determined for the seasonal component and merged with the smoothed data.

The second method used for trend analysis was a combination of the Mann-Kendall trend test (71, 72) and the Theil-Sen approach for slope estimation (73, 74). The Mann-Kendall test is a non-parametric test used to determine whether data have a significant monotonic upward or downward trend over time. The Theil-Sen approach is a non-parametric method of estimating the slope of a trend by determining the median slope through all possible pairs of points for the model predictions over time. However, neither the Theil-Sen approach nor the Mann-Kendall test is applicable to autocorrelated data such as environmental data with seasonal and inter-annual variation (75). Applying a pre-whitening procedure can remove autocorrelation, but at the risk of reducing the statistical significance and magnitude of real trends. Pre-whitening is a term used to describe the process of removing systematic noise from data to analyze underlying trends.

The approach proposed by S. Yue et al. (76) illustrates that the effect of removing autocorrelation on the estimate of the magnitude of the trend far exceeds the effect of removing the trend on the estimate of autocorrelation. Therefore, they propose the following method: 1) determine the slope of the trend using the Theil-Sen approach, and if the slope of the trend is significantly different from zero it is assumed to be linear and removed from the series; 2) calculate and remove the autocorrelation coefficient from the detrended series; 3) merge the detrended data with autocorrelation removed with the linear trend determined using the Theil-Sen approach; and 4) apply the Mann-Kendall test to the merged data to determine the significance of the trend. These methods were applied using the R statistical software package "*zyp*" (77). The slope of the trend was determined using monthly averaged data from 2020 to 2099 for all climate data and modeled lake groups to accommodate the computing power of the R statistical software. The long-term trend for each season was determined using daily data for each modeled lake group.

Lake model surface temperature predictions in 2020 and 2099 were analyzed to determine the change in the period when surface temperature was predicted to be above minimum and optimum growth temperatures for cyanobacteria as a function of climate, lake size, and trophic state. For each year, we determined the number of days that the surface water temperature was predicted to be above minimum growth temperatures, T_{min} (15°C) (14,15,) and the number of days the temperature is above optimal growth temperatures, T_{opt} (25°C) (14,15). The changes in the number of days above T_{min} and T_{opt} were compared across climates, lakes sizes, and trophic states to determine if any relationships existed and to determine the susceptibility of lakes to future cyanobacteria growth.

4. Results

4.1 Sensitivity Analysis

The sensitivity of the model's surface and mean water temperature predictions to modeled solar radiation, air temperature, air humidity, wind speed, cloudiness, and extinction coefficient were tested using observations at Allequash Lake (Table 3.5) to illustrate the effect of varying parameters by $\pm 10\%$ and $\pm 50\%$. The modeled surface (Figure 4.1) and mean (Figure 4.2) water temperature predictions appear to be most sensitive to solar radiation, air temperature, and air humidity. Although surface and mean water temperature predictions appear to be sensitive to the same parameters, modeled surface temperature predictions are much more variable than mean temperature predictions, which have a smoother response to varied climate inputs because of the lesser influence of the meteorological conditions at depth. There is also a seasonal component to the sensitivity of the lake model predictions for all varied parameters. In summer, the model is more sensitive to varied inputs than in winter months (Figure 4.1 and Figure 4.2).

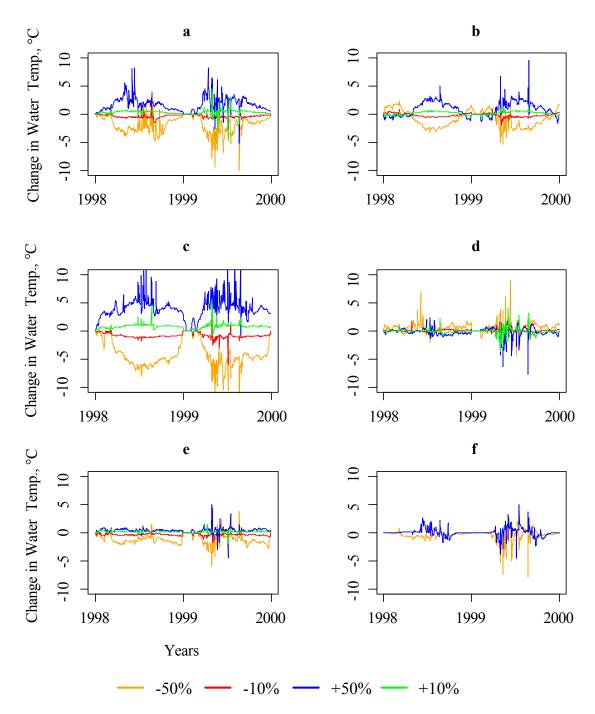


Figure 4.1. Response of surface water temperature predictions to varied modeled climate inputs and extinction coefficient including: a) solar radiation, b) air temperature, c) air humidity, d) wind speed, e) cloudiness, and f) extinction coefficient for Allequash Lake.

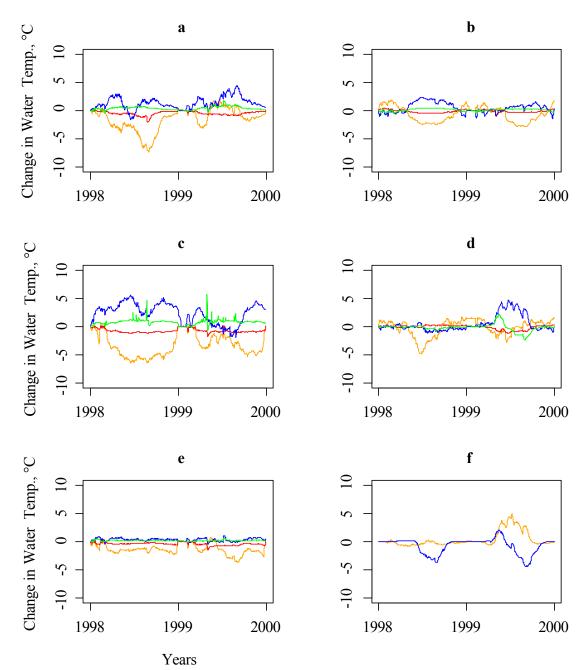


Figure 4.2. Response of mean water temperature predictions to varied modeled climate inputs and extinction coefficient including: a) solar radiation, b) air temperature, c) air humidity, d) wind speed, e) cloudiness, and f) extinction coefficient for Allequash Lake.

4.2 Modeled Meteorological Data

Plots of the difference between modeled and observed meteorological data from 1998 through 1999 illustrate the accuracy of the climate model (Figure 4.3). The errors were large for solar radiation, air temperature, and air humidity (Table 4.1). The differences between modeled and observed climate data were larger in warmer months for solar radiation and air humidity. The climate model deviation from observed conditions does not appear to vary much between the northern and southern validations sites.

Table 4.1. Climate model and observed means with root mean squared error (RMSE) for solar radiation, air temperature, air humidity, and wind speed for the northern and southern validation sites from 1998 through 1999.

Climate	Northern Validation Site			Southern Validation Site			
Parameter	Mean		DMCE	Me	DMCE		
	Observed	Modeled	RMSE	Observed	Modeled	RMSE	
Solar Radiation (W/m ²)	147.5	114.9	97.3	160.4	131.9	109.0	
Air Temperature (°C)	6.2	4.5	6.3	9.5	7.9	6.0	
Air Humidity (mb)	8.6	20.3	13.6	10.5	23.2	14.8	
Wind Speed (m/s)	2.6	4.0	2.6	3.3	4.4	2.8	

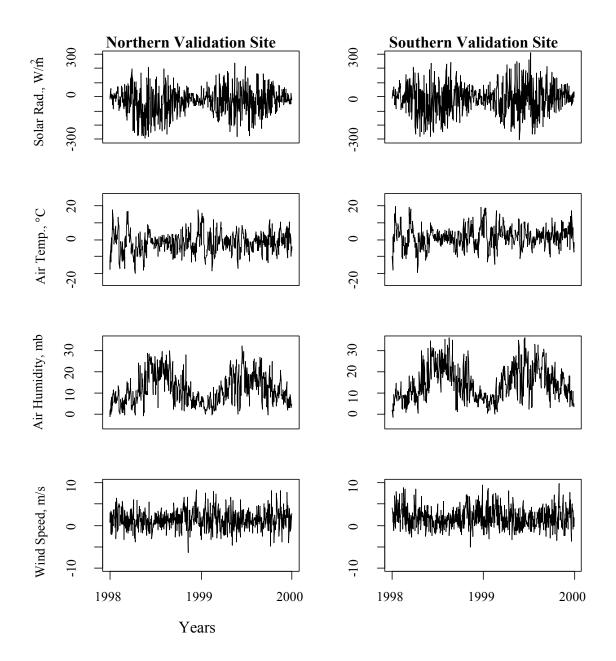


Figure 4.3. Difference between modeled and observed (modeled – observed) meteorological data for the northern (left column) and southern (right column) validation sites in Wisconsin from 1998 through 1999. Climate variables (solar radiation, air temperature, air humidity, and wind speed) are shown by row.

4.3 Validation

FLake's ability to predict surface and mean water temperatures was tested for 10 lakes. The model was found to be valid using the criteria described in section 3.5 and observed meteorological data for both mean and surface water temperature for Allequash

Lake, Crystal Lake, and Crystal Bog. The model predictions were valid for surface water temperature but not for mean water temperature for Big Muskellunge Lake, Sparkling Lake, Fish Lake, Lake Monona, and Lake Mendota. For Lake Wingra and Trout Lake, the model was not valid for either mean or surface water temperatures (Table 4.2). For supporting material see Appendix B.

The model validation was repeated using the modeled climate data for 1998 and 1999. Using the looser criteria, the lake model predictions were valid for predicting surface water temperature for all lakes except Crystal Bog. The model was found to be valid for mean water temperature for all lakes except Big Muskellunge Lake, Lake Wingra, Lake Monona, and Lake Mendota. The intercept showed that the lake model over-predicts when using modeled climate inputs for all lakes; the bias ranged from 3.4°C to 5.1°C. Results of all equivalence tests are shown in Table 4.2 and Appendix B.

T -l	Demonstern	Observed Climate Data	Modeled Climate Data	
Lake	Parameter	Valid	Valid	
		Y/N	Y/N	
A 11 1 T 1	Ts	Y	Y	
Allequash Lake	T _m	Y	Y	
Big Muskellunge	Ts	Y	Y	
Lake	T _m	Ν	Ν	
Crystal Lake	Ts	Y	Y	
	T _m	Y	Y	
Crystal Bog	Ts	Y	Ν	
	T _m	Y	Y	
Sparkling Lake	Ts	Y	Y	
	T _m	Ν	Y	
Trout Lake	Ts	Ν	Y	
	T _m	Ν	Y	
Fish Lake	Ts	Y	Y	
	T _m	Ν	Y	
Lake Wingra	Ts	Ν	Y	
	T_m	Ν	Ν	
Laka Manana	Ts	Y	Y	
Lake Monona	T_m	Ν	Ν	
Laka Mardata	Ts	Y	Y	
Lake Mendota	T_m	Ν	Ν	

Table 4.2. Summary of validation results for mean (T_m) and surface (T_s) water temperature using observed and modeled climate data. Validation entry results are shaded for cases that are not valid.

4.4 Trend Analysis

4.4.1 Seasonal Decomposition

Seasonal decomposition was first applied to all of the input climate variables for the Upper Peninsula, northern Lower Peninsula, and southern Lower Peninsula. Similar trends were observed for each variable among climates. The seasonally decomposed trends for solar radiation, air temperature, air humidity, wind speed, and cloudiness is illustrated using data from the northern Lower Peninsula in Figure 4.4. There appears to be a clear overall increasing trend in air temperature and air humidity while solar radiation, wind speed, and cloudiness seem to remain relatively constant over the modeled period (Figure 4.4) or decrease.

Seasonal decomposition was also applied to surface water temperature predictions for each of the modeled lake groups to illustrate long-term trends. The seasonally decomposed trends were similar across the modeled lake groups. Differences in the intercepts of the trends were observed among lakes in the Upper Peninsula, northern Lower Peninsula, and southern Lower Peninsula, while no differences were apparent among size or trophic state classes (Figure 4.5). Small, eutrophic lakes illustrate differences in trends among climate regions (Figure 4.5a). Eutrophic lakes in the southern Lower Peninsula illustrate differences among lake size classes (Figure 4.5b), and the results for large lakes in the Upper Peninsula illustrate differences among trophic states (Figure 4.5c).

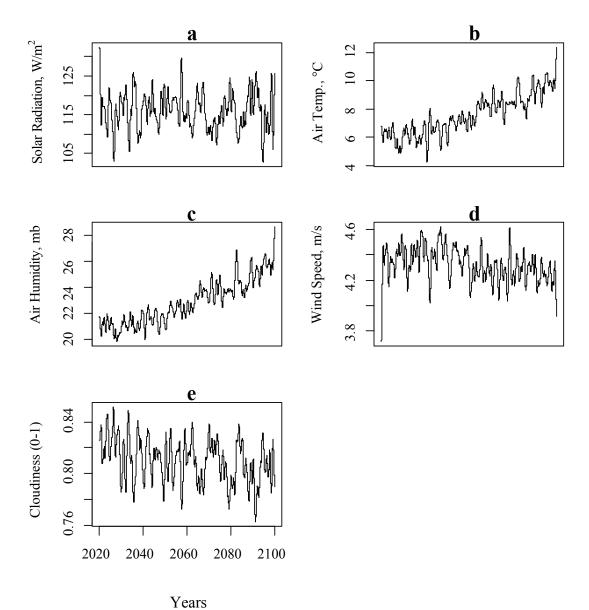


Figure 4.4 Seasonally decomposed trends in (a) solar radiation, (b) air temperature, (c) air humidity, (d) wind speed, (e) and cloudiness in the northern Lower Peninsula from 2020 to 2099.

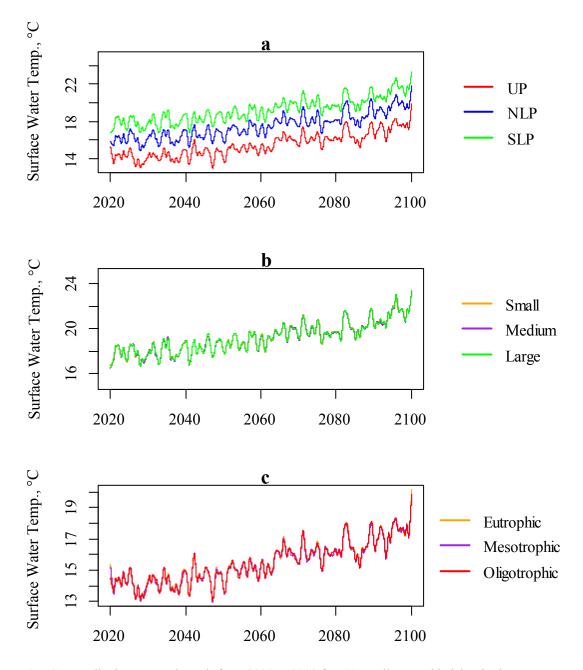


Figure 4.5. Seasonally decomposed trends from 2020 to 2099 for: (a) small, eutrophic lakes in the Upper Peninsula (UP-Sm-E), northern Lower Peninsula (NLP-Sm-E), and southern Lower Peninsula (SLP-Sm-E); (b) small, medium and large eutrophic lakes in the southern Lower Peninsula (SLP-Sm-E, SLP-Md-E, and SLP-Lg-E); (c) large eutrophic (UP-Lg-E), mesotrophic (UP-Lg-M), and oligotrophic lakes (UP-Lg-O) in the Upper Peninsula.

4.4.2 Slope Estimation

The Theil-Sen slope was estimated for each climate parameter in each climate region. Trends were significant for each of the following meteorological variables in all three climate regions: air temperature, air humidity, and cloudiness (Table 4.3). There was also a significant trend in wind speed in the Upper Peninsula (Table 4.3). The confidence intervals for the estimated trend in air temperature and humidity were compared between climate regions, and no statistical differences were observed. A summary of the mean long-term change in air temperature, air humidity, and cloudiness for all climate regions is reported in Table 4.4.

Table 4.3 Estimates of slope, 95% confidence intervals for the slope, intercept, and total change from 2020 to 2099 for each climate parameter in the Upper Peninsula (UP), northern Lower Peninsula (NLP), and southern Lower Peninsula (SLP). Significant trends are noted with an asterisk (*).

Location	Slope (^{unit} /decade)	CI_{Upper} (^{unit} /decade)	CI _{Lower} (^{unit} /decade)	Intercept (unit)	Total Change	
	Solar Radiation (W/m ²)					
UP	-0.2	0.1	-1.6	118.0	-1.3	
NLP	0.4	0.2	-1.1	123.4	3.3	
SLP	0.4	0.2	-1.2	130.6	3.4	
	A	ir Temperature	e (°C)			
UP	0.5	0.8	0.3	5.5	4.3*	
NLP	0.5	0.8	0.3	7.7	4.3*	
SLP	0.5	0.7	0.3	9.0	4.1*	
		Air Humidity (mb)			
UP	0.6	0.8	0.4	19.1	4.8*	
NLP	0.7	0.9	0.4	21.0	5.4*	
SLP	0.7	1.0	0.4	23.4	5.8*	
		Wind Speed (r	m/s)			
UP	0.0	0.0	-0.1	4.5	-0.2*	
NLP	0.0	0.0	0.0	4.5	-0.1	
SLP	0.0	0.0	-0.1	4.8	-0.2	
		Cloudiness (%)			
UP	-0.3	-0.1	-0.5	0.8	-2.2*	
NLP	-0.4	-0.2	-0.7	0.8	-3.5*	
SLP	-0.4	-0.1	-0.6	0.8	-3.0*	

Parameter	Mean Change from 2020 to 2099
Solar Radiation (W/m ²)	1.8
Air Temperature (°C)	4.3*
Air Humidity (mb)	5.3*
Wind Speed (m/s)	-0.2
Cloudiness (%)	-2.9*

Table 4.4. Mean change in air temperature, air humidity, and cloudiness between 2020 and 2099 for all climate regions. An asterisk (*) identifies trends that were statistically significant.

The annual and seasonal trends in surface water temperature from 2020 to 2099 were significant for all modeled lake groups (Table 4.5). For all lake groups, the largest seasonal increase in surface water temperature was observed in spring and the smallest increase was in winter. No statistical differences in long-term trends of surface temperature were identified among any of the groups regardless of location, size, or trophic state (Table 4.5). However, across all lake groups, the long-term trend in winter was different from spring, summer, and fall while the long-term trend in fall was not statistically different from summer (Table 4.6). The long-term trend in fall was different from spring with the exception of three lake groups: small eutrophic and mesotrophic lakes and medium eutrophic lakes in the southern Lower Peninsula (SLP-Sm-E, SLP-Sm-M, and SLP-Md-E, respectively), but there was no obvious relationship between this result and climate, lake size, or trophic state. Fall increases in surface temperature were greater than summer increases in all groups except: small, eutrophic lakes in the Upper Peninsula (UP-Sm-E); large eutrophic, mesotrophic, and oligotrophic lakes in the Upper Peninsula (UP-Lg-E, UP-Lg-M, and UP-Lg-O, respectively); and large oligotrophic lakes in the northern Lower Peninsula (NLP-Lg-O).

Table 4.5 Estimates of slope, 95% confidence intervals for the slope, intercept, and total change in surface water temperature from 2020 to 2099 are reported for each modeled lake group for the entire period. All trends (slopes) were significant at $\alpha = 0.05$.

Crown	Slope	CI _{Upper}	CI _{Lower}	Intercept	Total
Group	(°C/decade)	(°C/decade)	(°C/decade)	(°C)	Change (°C)
UP-Sm-E	0.5	0.7	0.3	13.9	3.8
UP-Md-E	0.5	0.7	0.3	13.8	3.8
UP-Lg-E	0.5	0.7	0.3	13.6	3.8
UP-Lg-M	0.5	0.7	0.3	14.0	3.8
UP-Lg-O	0.5	0.7	0.3	13.8	3.8
NLP-Sm-E	0.5	0.7	0.3	15.6	4.0
NLP-Lg-O	0.5	0.7	0.3	16.0	3.9
SLP-Sm-E	0.5	0.7	0.3	17.6	3.8
SLP-Sm-M	0.5	0.7	0.3	17.5	3.8
SLP-Sm-O	0.5	0.7	0.2	17.8	3.8
SLP-Md-E	0.5	0.7	0.3	17.6	3.8
SLP-Lg-M	0.5	0.7	0.3	17.5	3.9
SLP-Lg-O	0.5	0.7	0.3	18.2	3.9

Group	Season	Slope (°C/decade)	CI _{Upper} (°C/decade)	CI _{Lower} (°C/decade)	Intercept (°C)	Total Change (°C)
	Fall	0.05	0.06	0.05	11.0	3.2
UP-Sm-E	Spring	0.08	0.09	0.07	15.3	4.9
UP-SIII-E	Summer	0.06	0.06	0.05	25.7	3.6
	Winter	0.03	0.03	0.02	2.5	1.6
	Fall	0.06	0.07	0.05	10.9	3.6
UP-Md-E	Spring	0.08	0.09	0.07	15.7	4.7
UP-Mu-E	Summer	0.06	0.06	0.05	25.5	3.6
	Winter	0.03	0.03	0.03	2.2	1.7
	Fall	0.06	0.06	0.05	10.9	3.4
LID L ~ E	Spring	0.08	0.09	0.07	15.7	4.7
UP-Lg-E	Summer	0.06	0.06	0.05	25.6	3.6
	Winter	0.03	0.03	0.03	2.3	1.7
	Fall	0.05	0.06	0.04	11.3	3.2
UD L ~ M	Spring	0.08	0.09	0.07	14.8	5.0
UP-Lg-M	Summer	0.06	0.09	0.05	25.5	3.6
	Winter	0.03	0.03	0.02	2.6	1.5
	Fall	0.05	0.06	0.05	12.6	3.2
	Spring	0.09	0.09	0.08	13.4	5.3
UP-Lg-O	Summer	0.06	0.06	0.05	25.1	3.6
	Winter	0.02	0.03	0.02	3.1	1.3
	Fall	0.06	0.07	0.05	13.3	3.5
	Spring	0.08	0.08	0.07	17.5	5.0
NLP-Sm-E	Summer	0.06	0.06	0.05	27.1	3.6
	Winter	0.03	0.04	0.03	4.2	2.0
	Fall	0.05	0.06	0.05	15.0	3.2
	Spring	0.09	0.10	0.09	15.5	5.7
NLP-Lg-O	Summer	0.06	0.06	0.05	26.9	3.5
	Winter	0.02	0.03	0.02	5.2	1.4
	Fall	0.06	0.07	0.06	14.7	3.8
	Spring	0.08	0.09	0.07	19.3	4.7
SLP-Sm-E	Summer	0.05	0.06	0.05	28.6	3.4
	Winter	0.03	0.03	0.02	6.7	1.7
	Fall	0.06	0.07	0.05	14.7	3.7
	Spring	0.08	0.08	0.07	19.3	4.7
SLP-Sm-M	Summer	0.05	0.06	0.05	28.5	3.4
	Winter	0.03	0.03	0.02	6.7	1.6

Table 4.6 Estimates of slope, 95% confidence intervals for the slope, intercept, and change in surface water temperature from 2020 to 2099 are reported for each modeled lake group for each season. All trends were significant at $\alpha = 0.05$. (Table continued on the next page.)

Group	Season	Slope (°C/decade)	CI _{Upper} (°C/decade)	CI _{Lower} (°C/decade)	Intercept (°C)	Total Change (°C)
	Fall	0.06	0.07	0.05	15.5	3.6
SLP-Sm-O	Spring	0.08	0.09	0.08	18.6	5.1
5LF-5111-0	Summer	0.05	0.05	0.05	28.4	3.2
	Winter	0.03	0.03	0.02	6.9	1.6
	Fall	0.06	0.07	0.05	14.4	3.7
SLP-Md-E	Spring	0.08	0.08	0.07	19.5	4.7
SLF-IVIU-E	Summer	0.06	0.06	0.05	28.7	3.4
	Winter	0.03	0.03	0.03	6.5	1.8
	Fall	0.06	0.07	0.05	14.5	3.5
SIDIAM	Spring	0.08	0.09	0.07	19.4	4.8
SLP-Lg-M	Summer	0.06	0.06	0.05	28.7	3.5
	Winter	0.03	0.03	0.03	6.5	1.8
	Fall	0.06	0.06	0.05	16.7	3.4
SLP-Lg-O	Spring	0.09	0.10	0.08	17.5	5.5
	Summer	0.05	0.06	0.05	28.1	3.3
	Winter	0.02	0.02	0.02	7.4	1.2

Table 4.7. Summary of mean change in surface water temperature for all modeled lake groups from 2020 to 2099.

Season	Mean Change in Surface Water Temperature (°C)
Fall	3.5
Spring	5.0
Summer	3.5
Winter	1.6
All	3.8

4.4.3 Cyanobacteria Growing Period

The change in the number of days the surface water temperature was above minimum and optimal temperatures for cyanobacteria growth were compared for the years 2020 and 2099 for each modeled lake group (Figure 4.6). Climate region had the greatest effect on changes in periods of both minimum and optimum growth temperatures. Periods of growth increased as latitude decreased for minimum growth temperatures and increased as latitude increased for optimum growth temperatures.

The increase in the period when temperatures were above minimum growth conditions was greatest for medium lakes, closely followed by large lakes, and small lakes showed the lowest change (Figure 4.6). The change in period when temperature was above optimum conditions as a function of lake size was not consistent among climates and trophic states. Medium lakes showed the greatest increase, closely followed by large lakes, and small lakes showed the lowest change for eutrophic lakes in the Upper

Peninsula, while the change in the period decreased with lake size for eutrophic lakes in the southern Lower Peninsula (Figure 4.6). The change in period when temperature was above minimum conditions across lake trophic state was affected by climate region and lake size. Mesotrophic lakes showed the lowest increase, while eutrophic and oligotrophic lakes were higher for large lakes in the Upper Peninsula (Figure 4.6). This result was inversed for small lakes in the southern Lower Peninsula (Figure 4.6). The change in period when temperature was above optimum conditions across lake trophic state was also dependent on climate and lake size. For large lakes in the Upper Peninsula, the increase in the period increases with water clarity, while for small lakes in the southern Lower Peninsula the greatest change was observed for oligotrophic lakes, followed by eutrophic and then mesotrophic lakes (Figure 4.6).

Analyzing trends between lake properties using the data for all modeled lake groups revealed interactive effects on the change in the period when temperature is above minimum and optimal growth temperatures. Using the data for all modeled lake groups, an increasing trend was observed in the days when temperature were above minimum growth temperatures increases as latitude and size decreases (Figure 4.7a,b). An increasing trend was observed in the days when temperatures were above optimum growth temperatures with decreasing latitude and increasing lake size (Figure 4.8a,b). In all cases the minimum and optimum temperature periods increased with increasing water clarity (Figure 4.7c and Figure 4.8c).

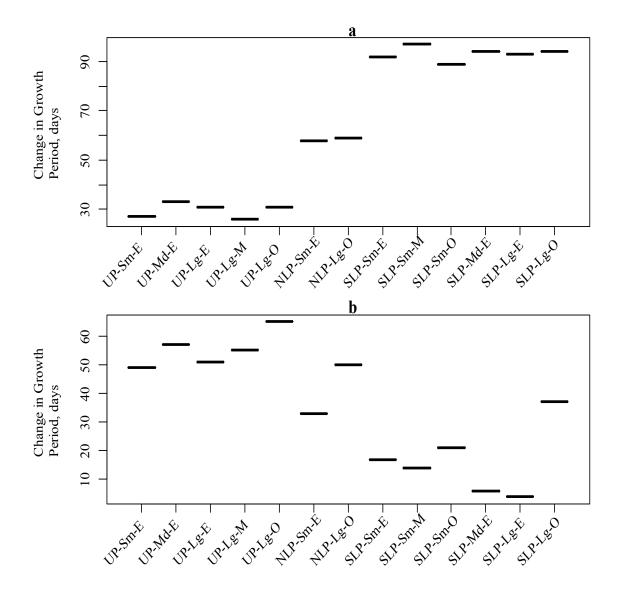


Figure 4.6. Summary of changes in number of days with surface temperature above a) T_{min} (15°C) and b) T_{opt} (25°C) from 2020 to 2099 for each modeled lake group.

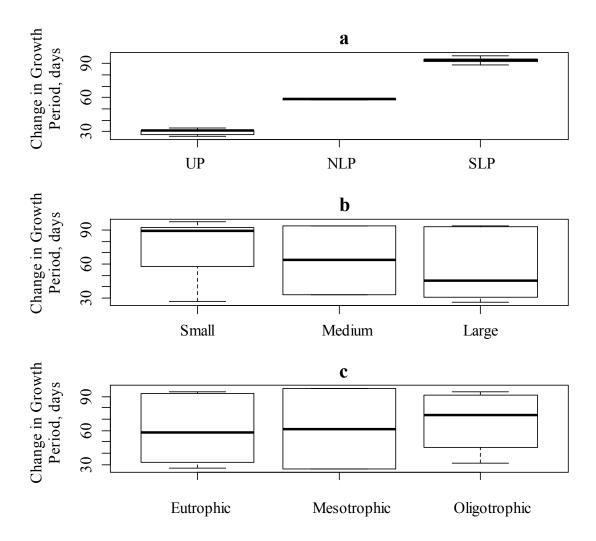


Figure 4.7. Change in the number of days when surface water temperature is above T_{min} (15°C) between 2020 and 2099 as a function of a) climate, b) lake size, and c) trophic state.

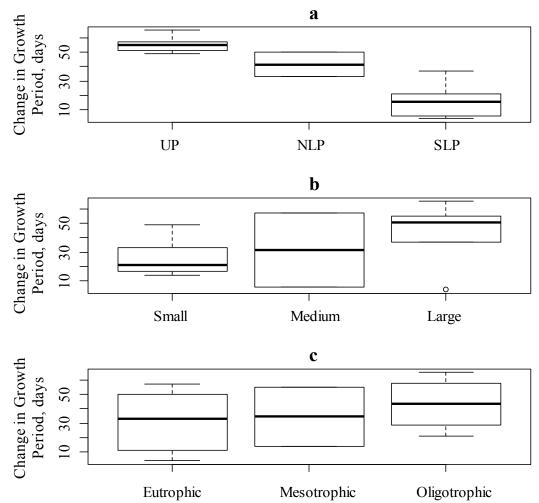


Figure 4.8. Change in the number of days when surface water temperature is above T_{opt} (25°C) between 2020 and 2099 as a function of a) climate, b) lake size, and c) trophic state.

5. Discussion

5.1 Sensitivity Analysis

The sensitivity analysis aided us in identifying the relative effects of meteorological drivers and water transparency on the model predictions. The model predictions showed the largest response to changes in air humidity, and the response of surface water temperature predictions is much more variable than the response of mean temperature predictions. This result is likely due to the exposure of the surface of the lake to the atmosphere, resulting in a more immediate response to external conditions compared to the rest of the water column. The method used to conduct the sensitivity analysis may have resulted in a misrepresentation of the relative effects of meteorological inputs on lake model predictions. Inputs were varied one at a time, but in the natural world this would not be the case. Most climate parameters do not vary independently and have a

combined effect on the lake heat budget. Each of the climate input parameters were varied by $\pm 10\%$ and $\pm 50\%$, but the amount of natural variability that these changes comprise is different for each climate variable (Table 3.6). The sensitivity analysis showed that the model is most sensitive to air humidity inputs, which would mainly affect latent heat fluxes. Latent heat fluxes do affect surface heat exchange, but are typically relatively small factors in the heat budget in comparison to sensible heat and radiation fluxes. The relative magnitude of the effect may have been skewed because input parameters were not co-varied.

5.2 Modeled Meteorological Data

The ability of the dynamically downscaled global climate model used in this study to reproduce observed conditions during the two year validation period was limited. The variability of model error was large for all climate variables and is not constant across seasons for solar radiation and air humidity. The climate model error appears to be relatively constant from year to year in the two years of simulation; however, there were not enough data to make a definitive conclusion.

5.3 Validation

The results of the validation of FLake using observed meteorological data showed that the model performed well for predicting surface water temperature. Reasons for the poorer fit for Lake Wingra and Trout Lake using observed climate data inputs may be that these lakes deviate from the model assumptions or other contributing factors were not represented in the model. For example, approximately 75% of the Lake Wingra watershed is urbanized and receives large amounts of phosphorus through surface runoff (78) and Trout Lake receives significant ground-water exchanges (79).

The lake model was valid for predicting mean water temperatures using observed meteorological data for only 3 out of 10 lakes. Based on this result, we do not recommend using FLake to predict mean water temperatures in absolute terms without lake-specific calibration. While being unable to model mean temperature limits the applicability of the model for determining climate impacts on aquatic organisms, this guidance is in line with the model developers' focus on surface water temperature (*33*).

The lake model was valid using modeled meteorological data to predict surface water temperature (9 of 10 lakes) but not mean temperature (6 of 10 lakes). These results were similar to the validation results using observed meteorological data, and indicate that modeled meteorological inputs allow FLake to predict surface water temperatures with reasonable accuracy. However, the looser validation criteria for the modeled climate data allowed for increased bias (greater than 1°C) in FLake predictions.

The response of the model predictions to humidity, coupled with the relatively high errors in humidity between the modeled and observed climate data indicate that the bias in the model predictions was mainly caused by inaccurate humidity predictions from the climate model. A $\pm 10\%$ change in air humidity altered surface and mean water temperature predictions by 2°C (Figure 4.1 and Figure 4.2). A 10% increase in air humidity increased the mean air humidity by 2 mb, and a 50% increase in air humidity increased the mean air humidity by 10 mb (Table 3.5). Climate model humidity predictions exceeded observations by up to 30 mb during the summer months (Figure 4.3) which may have drastically altered the latent heat flux. Although the latent heat flux

is a relatively small component of surface heat exchange, increased air humidity can favor condensation instead of evaporation, and increase the heat flux into the lake. While air humidity was exaggerated by the inaccurate climate model predictions, it is likely that the bias was also linked to error in other climate model parameters such as air temperature and solar radiation.

5.4 Trend Analysis

A strong increasing trend in surface water temperature was observed from 2020 to 2099 for all modeled lake groups. The seasonally decomposed trend in surface water temperature for all modeled groups closely followed trends in air temperature. This is likely because air temperature plays a key role in lake surface heat exchange through sensible and latent heat fluxes. There were no observable differences in the trends in lake surface temperature for different lake sizes or trophic states, but there was an observable difference in the intercept among climate groups. Because there were no statistical differences in the modeled climates used to drive the lake model predictions, the differences in the intercepts of surface temperature trends may have been a function of the latitudes used as model inputs, and the sensitivity of the model to latitude should be tested. Our results agree with findings from studies using both contemporary evidence and lake models (28,28,35). Similar studies found that surface water temperature trends closely followed local air temperatures with no detectable differences in trends across lake morphologies or trophic states (28,28,35). Additionally, maximum surface water temperatures were found to be most affected by air temperature, dew point temperature, and solar radiation (28).

The results of this study were surprising given our knowledge that surface heat exchange is controlled, in part, by lake surface area (18, 19) and that water clarity can affect radiation reflection in the lake (21). Similarly, location affects atmospheric conditions and we expected to see differences in climate and surface water temperature trends among the three climate regions. Our main explanation for the lack of differences in surface temperature responses among climates was that there were no difference in air temperature trends among the climate regions. It is unclear if the lack of difference in air temperature trends was real or a product of model error. The Great Lakes act as a buffer for local air temperatures and reduce the regional temperature range (80, 81), and so air temperatures may not actually vary much throughout Michigan's Upper and Lower Peninsula. However, the results of the sensitivity analysis showed that the climate model has considerable error, which may have muted the differences in temperature among the climate regions.

5.5 Ecological Implications

The results of the slope estimation showed that the greatest predicted increase in surface water temperature occurred in spring, followed by summer and fall which resulted in longer periods of minimum and optimum growth conditions for cyanobacteria. Change in the period when the surface water temperature was predicted above minimum and optimum growing conditions between 2020 and 2099 varied by climate, lake size, and trophic state. Using the data for all modeled lake groups, the number of days when lake temperatures were above minimum growth temperatures increased as latitude and size decreased, and increased with increasing water clarity. Lakes in the southern Lower

Peninsula are subject to warmer temperatures and increased solar radiation earlier in the year than more northern climate regions, allowing lakes to reach minimum growth temperatures earlier in the spring. Furthermore, small lakes are able to warm more quickly than large lakes. The link between longer periods above minimum growth temperature in oligotrophic lakes was likely due to the higher transmission through and absorption of solar radiation in the clearer waters. Based on these results, it follows that small, oligotrophic lakes in the southern Lower Peninsula will be most at risk for cyanobacteria blooms because of the large number of days above the minimum water temperature.

The change in number of days when temperatures were above optimum growth temperature decreased with decreasing latitude and increased with increasing size and water clarity. The effect of climate was unexpected, as it would seem that the effect of climate on sustained minimum and optimum temperatures would be the same. However, because we are considering the change in the growth period, this result reflects large number of days when surface temperatures are already above optimum growth temperatures in the more southern regions. The increased period for optimal growth temperatures in the colder Upper Peninsula was greater because of the relatively low number of days at the onset of the modeling period in comparison to more southern climate regions. The implication is that lakes in the Upper Peninsula will experience a much greater change in risk of cyanobacteria blooms than present, while the risk will be only slightly elevated in the southern part of the state.

The increasing effect on the change in the period supporting optimum growth conditions with increasing lake size is more straightforward. Large lakes tend to uptake more heat and remain stratified for longer into the summer than small lakes. Although they do not meet minimum growth temperatures as quickly as small lakes, large lakes are able to absorb more heat and release heat a slower rate as summer approaches fall resulting in longer periods of time when temperatures are above optimum growing conditions than small lakes.

Validation using modeled climate data showed that lake model surface temperatures were over-predicted by an average of approximately 4°C. This likely resulted in an overestimate of the change in the time period for minimum and optimum growth conditions. However, our results were used to indicate relative importance of climate, size, and trophic state, despite modeling errors that limit the use of the absolute periods within minimum and optimal growing temperatures.

5.6 Limitations and Future Work

The results of this study are limited by a number of factors that should be considered when interpreting the outcomes. These limitations also present opportunities for future work that will improve and expand our results. First, we must be mindful of our current ability to model lakes and climate. We demonstrated the accuracy of a regionally downscaled global climate model and a lake model, as well as the impact of accuracy of the climate model on the accuracy of the lake model predictions. The effect of the variability and magnitude of climate model error on error in the lake model predictions is apparent from the validation and sensitivity analysis results. As downscaling methods for regional climate models are further developed, these methods can be reapplied to increase the certainty of the lake model predictions. With improved climate and lake model predictions, these results may be strengthened. It may also produce the ability to detect differences between lake types not detected herein and therefore help refine management practices for the future.

The lake model used in this study also has limitations. This study attempted to determine the response of a variety of lake types with a very limited amount of observed data for validation and model inputs. Because of the relatively high error rates during validation, we chose not to use FLake to predict mean water temperatures under future conditions. A study with the primary focus of validating FLake using lakes with a variety of physical and chemical properties would help to better define the limits and applicability of the model in terms of lake characteristics. Future work quantifying the uncertainty from the coupled climate and lake model would be another way to improve our analysis and strengthen the results using methods such as Monte Carlo simulations *(82)* that capture concurrent variability over time.

FLake is a simplified bulk model that is meant to be used without lake-specific calibration, which limits the user's ability to adjust the model parameters to achieve better results. One option for improving lake model results would be to identify the bounds of the model's applicability for different lake types and alter the sampling scheme to accommodate those results. Another potential avenue for model improvement may be to expand the self-similarity concept applied to the thermocline and hypolimnion to include other possible shapes of the temperature-depth curve. Other options to improve the accuracy of future predictions would be to choose a lake model that better captures lake processes with more complexity, such as one-dimensional finite-difference models (37,38,39,40) or a three-dimensional hydrodynamic model (83).

Using prototype lakes to represent lake groups allowed us to extrapolate our results to a larger number of lakes. This, in turn, would make the results more applicable for management applications as they are not for specific lakes, but categories of lakes with similar characteristics. In contrast, using prototype lakes eliminated our ability to calibrate the model and validate the accuracy of model predictions. Other disadvantages included the bias in the lake size distribution of the MSU dataset that may have resulted in an overestimate of the size of the prototype lake size for small lakes and an underestimate in the size of large lakes in comparison to the actual lake population in Michigan. Additionally, the small lakes in the MSU dataset tend to have steeper basin slopes than large lakes, creating a bias in the morphology of the prototype lakes.

The effects of increasing surface water temperatures have a much larger scope than what was presented here, which focused on effects on algal growth. Increasing surface water temperatures have implications for ice duration, mixing regime, duration of stratification, and food web structure that which can affect the ecology in a lake. The results of this study lay the groundwork for further investigation into effects of climate change on a large number of lake properties, including hypolimnetic temperatures, mixing, ice cover, and lake ecology. Future work on these topics will help us better characterize lakes' physical responses to climate change and subsequent ecological effects.

6. Conclusions

A one-dimensional, bulk lake model was selected and validated to model the impacts of climate change on the surface water temperature of Michigan's inland lakes from 2020 to 2099. Lake model predictions were analyzed for long-term trends, and potential effects on cyanobacteria growth were identified. The lake model FLake accurately reproduced surface water temperatures without calibration; however, it did not pass validation criteria for predicting mean water temperature using either observed or modeled climate data. The comparison of modeled and observed climate data showed that the regionally downscaled global climate model used in this study produced highly variable error. We found that FLake was especially sensitive to solar radiation, air temperature, and air humidity and this highlighted the importance of accurate climate model data for accurate lake model predictions.

Seasonal decomposition and slope estimation showed long-term trends for each of the climate parameters and for lake surface temperature predictions. Positive trends in air temperature and humidity were significant, as was a slightly decreasing trend in cloud cover. Long-term trends in surface water temperature were significant for all groups, and there were no statistical differences in slopes among the modeled lake groups. The intercepts of the trends among climate groups increased as the latitude decreased. Significant differences in long-term trends were identified among seasons, and for all modeled lake groups the long-term trend from 2020 to 2099 was most gradual in winter and steepest in spring.

Analysis of the change in the period of minimum and optimal growth temperatures for cyanobacteria between 2020 and 2099 showed that climate region will play a key role in determining the risk of future cyanobacteria growth. The predicted change in the number of days when surface water temperature was above the minimum growth temperature increased as latitude decreased. Conversely, the predicted change in the number of days when surface water temperature was predicted to be above optimum growth temperatures increased as latitude increased. When considering the effects of lake size and trophic state, the largest increase in the period with surface temperatures above minimum growth temperature was predicted for small, oligotrophic lakes in the southern Lower Peninsula. The largest increase in the period with surface temperatures above optimum growth temperature was predicted for large, oligotrophic lakes in the Upper Peninsula.

Through this study we provided an approach for large scale lake studies, quantitative validation, and trend analysis. This study also supports a more holistic understanding of climate change effects on small lakes and the results can be used to inform the development of plans to prevent or mitigate algal growth in lakes that will be at greater risk in future climates.

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Appendices

A. Lakes and Classifications

Table A.1 reports the lake name, surface area, secchi depth, maximum depth, and mean depth for each of the 517 lakes in the study area (47). The lake data are reported by lake group. There are 27 lake groups based on climate, size, and trophic state. An entry of 'NA' indicates that there was no data available.

Name	Surface Area (km ²)	Secchi Depth (m)	Max Depth (m)	Mean Depth (m)
Group	l: Upper Peninsu	lla, Small, Eutro	phic	· · ·
Sunken Lake	0.30	1.83	6.40	3.40
Bass Lake	0.41	0.76	6.40	2.10
Six mile Lake	0.40	1.68	7.01	3.00
Pickerel Lake	0.28	1.98	18.59	5.50
Pole Creek Lake	0.36	0.61	3.05	1.50
Dana Lake	0.34	1.68	7.62	NA
Ashford Lake	0.06	1.37	12.19	NA
Trout Lake	0.24	0.76	2.13	NA
Little Oxbow Lake	0.39	1.52	18.29	4.90
Taylor Lake	0.45	1.37	12.19	5.10
Boot Lake	0.43	1.76	9.14	3.60
Kaks Lake	0.24	1.52	6.71	2.50
Culhane Lake	0.42	1.37	14.94	4.50
Pike Lake	0.33	1.52	15.55	3.30
Group 2:Upper Penin	sula, Small, Meso	otrophic		
Crooked Lake	0.39	3.20	8.53	3.80
Pickerel Lake	0.38	3.66	21.95	7.80
KP Lake	0.44	3.66	7.62	3.50
Section One Lake	0.24	2.74	8.23	2.50
Bass Lake	0.19	2.74	6.10	NA
Emerald Lake	0.25	3.35	10.67	3.00
Gaylanta Lake	0.45	3.05	15.24	3.50
Dixon Lake	0.32	3.81	9.14	4.10
Lake Twentyseven	0.43	3.20	6.71	3.60
Hoffman Lake	0.48	3.96	9.14	2.60
Tomahawk Lake	0.17	3.02	9.75	NA
Shoepac Lake	0.21	2.74	28.65	NA
Lost Lake	0.42	3.35	5.18	1.50

Table A.1. Summary of lake property data for the 545 lakes in the study by lake group.

Name	Surface Area (km ²)	Secchi Depth (m)	Max Depth (m)	Mean Depth (m)
Lancaster Lake	0.21	3.35	17.37	8.20
Mary Lake	0.35	3.35	25.30	6.50
Hamilton Lake	0.31	3.96	9.14	3.30
Rock Lake	0.35	2.29	9.14	3.60
Carney Lake	0.46	3.05	10.67	4.00
Camp Seven Lake	0.21	3.81	10.67	5.00
Little Smoky Lake	0.35	3.05	6.10	3.10
Unnamed	0.35	2.44	15.24	3.90
Allen Lake	0.32	3.05	7.62	4.50
Gibson Lake	0.37	2.44	7.01	2.70
Anderson Lake	0.21	3.96	9.14	2.40
Pike Lake	0.37	3.96	9.14	3.70
Moraine Lake	0.37	3.96	6.10	2.10
Eel Lake	0.21	3.66	7.62	2.70
Island Lake	0.09	2.01	12.19	NA
County Line Lake	0.28	2.90	13.72	5.40
Deer Lake	0.32	3.20	13.41	4.50
Sporley Lake	0.31	3.35	12.80	6.40
Bobcat Lake	0.36	2.13	5.49	2.00
Hannah Webb Lake	0.26	3.81	10.67	4.00
Snyder Lake	0.25	3.20	5.18	1.90
McClure Storage Basin	0.42	2.44	14.60	4.60
Wolf Lake	0.46	2.44	3.96	1.60
Perch Lake	0.38	2.44	15.24	5.60
Muskrat Lakes	0.04	2.44	7.93	NA
Unnamed	0.06	3.05	7.32	NA
Group 3:Upper Penins	ula, Small, Olig	otrophic		
Cub Lake	0.23	5.18	7.01	2.90
Shupac Lake	0.43	7.92	29.57	8.70
Heart Lake	0.27	4.72	35.66	9.80
Big Bass Lake	0.26	5.79	10.06	3.50
Lake Fifteen	0.37	4.27	16.76	9.40
Ess Lake	0.48	5.18	15.55	5.10
Stager Lake	0.44	5.18	16.76	5.80
Camp Lake	0.40	4.27	17.07	5.50
Unnamed	0.22	4.88	7.62	4.10
Long Lake	0.24	6.10	32.00	9.60
Dinner Lake	0.44	4.02	12.19	3.80
Imp Lake	0.37	5.94	26.21	9.60

Name	Surface Area (km ²)	Secchi Depth (m)	Max Depth	Mean Depth
Edey Lake	0.33	4.88	(m) 9.14	(m) 2.90
Silver Lake	0.44	4.88	7.01	2.90
Fish Lake	0.49	4.27	12.19	5.00
Unnamed	0.21	4.27	19.81	4.60
Bass Lake	0.31	7.01	23.77	9.90
Little Shag Lake	0.43	4.88	10.67	5.50
Johnson Lake	0.32	4.27	13.72	5.90
Unnamed	0.50	4.27	10.67	2.50
Norway Lake	0.21	5.99	6.10	2.60
Big Lake	0.48	4.72	10.36	1.30
Lake Arfelin	0.23	4.88	10.67	5.10
Emily Lake	0.24	4.27	27.43	7.40
Group 4:Upper Penins				
Lake May	0.77	1.83	6.10	2.30
Brule Lake	0.98	0.91	6.10	3.30
Paint Lake	0.96	1.68	4.57	NA
Winslow Lake	1.05	1.98	6.10	2.80
Ruth Lake	0.81	1.83	10.97	3.40
Bob Lake	0.53	1.26	4.57	2.30
Lake Roland	1.05	1.83	12.19	4.10
Thayer Lake	0.52	1.52	3.05	1.00
Group 5:Upper Penins	ula, Medium, M	lesotrophic		
Lake Manuka	0.64	2.74	8.23	1.60
Big Lake	0.50	3.78	24.69	7.30
Unnamed	1.13	3.96	24.38	6.80
Unnamed	0.92	2.13	4.57	NA
Lake Emma	0.80	3.05	3.05	1.00
Buck Lake	0.61	2.74	9.14	3.40
Lake Mary	1.06	3.91	14.63	6.30
Swan Lake	0.65	2.44	6.10	4.40
Frenchman Lake	0.75	2.13	6.10	2.10
James Lake	0.83	2.74	3.05	1.60
Fire Lake	0.51	2.74	12.19	3.30
Petes Lake	0.78	3.81	11.28	4.10
Moosehead Lake	0.58	2.74	11.89	4.60
Clearwater Lake	0.72	2.44	3.05	2.10
Marion Lake	1.20	3.35	12.19	3.70
Shag Lake	0.79	3.96	9.10	4.60
Bass Lake	0.80	2.59	5.49	2.60

Name	Surface Area (km ²)	Secchi Depth (m)	Max Depth (m)	Mean Depth (m)
Horseshoe Lake	0.53	2.74	7.32	2.60
Unnamed	0.83	3.96	28.96	7.30
Lake Sainte Kathryn	0.67	2.29	7.62	3.10
Deer Lake	1.07	2.92	21.95	NA
Monocle Lake	0.69	2.74	16.76	6.50
Gemini Lakes	0.51	2.59	6.10	2.70
Ross Lake	0.79	2.77	6.10	3.00
Lake Keewaydin	0.53	2.29	7.60	2.90
Kingston Lake	0.51	2.90	5.50	NA
Pike Lake	1.17	2.13	13.11	4.70
Bodi Lake	1.19	2.13	14.63	3.50
Lake Fanny Hooe	0.93	3.96	12.19	6.40
Group 6:Upper Penins	ıla, Medium, O	ligotrophic		
	Surface Area	Secchi Depth	Max	Mean
Name	(sq. km)	(m)	Depth	Depth
			(m)	(m)
Opal Lake	0.51	5.18	12.19	4.60
Avery Lake	1.18	7.01	27.43	8.90
Lake Nettie	1.20	4.53	14.02	3.70
Indian Lake	0.80	5.03	10.97	4.60
Island Lake	0.76	5.50	17.53	5.90
Golden Lake	1.10	8.65	30.48	8.00
Lake Ellen	0.57	4.57	18.29	6.90
Colwell Lake	0.52	4.72	7.62	4.60
Unnamed	1.16	5.49	24.38	8.20
Bass Lake	0.60	4.11	22.56	8.30
Group 7:Upper Penins	ıla, Large, Eutr	ophic		
Grand Lake	23.57	1.81	7.62	2.70
Alke Paradise	7.76	1.51	5.18	NA
Unnamed	2.81	0.91	9.40	3.00
Tamarack Lake	1.34	1.22	4.57	3.48
Langford Lake	1.95	1.51	3.05	2.10
Chaney Lake	2.01	1.83	6.10	2.20
McDonald Lake	1.87	1.22	3.05	1.00
Nawakwa Lake	1.79	1.37	10.67	3.00
Prickett Lake	3.11	1.22	17.07	5.67
Otter Lake	3.64	1.37	8.84	4.80
Group 8:Upper Penins	ıla, Large, Mes	otrophic		
Lake Margrethe	7.79	3.51	19.81	4.70

Name	Surface Area (km ²)	Secchi Depth (m)	Max Depth (m)	Mean Depth (m)
Manistee Lake	3.54	2.77	5.49	2.20
East Twin Lake	3.32	2.13	7.62	2.00
Otsego Lake	8.15	3.20	7.01	3.00
Beaver Lake	2.74	3.51	23.47	8.50
Rush Lake	1.72	3.05	10.10	2.20
Long Lake	21.64	2.51	7.62	3.20
Gulliver Lake	3.56	2.26	8.53	3.50
Stanley Lake	1.29	3.05	11.89	3.30
Iron Lake	1.58	2.13	13.72	3.00
Sunset Lake	2.15	3.66	16.46	5.00
Lake Emily	1.32	2.44	9.75	3.40
Round Lake	1.96	3.96	17.07	5.80
Duck Lake	2.49	2.90	7.60	4.10
Little African Lake	6.81	2.51	9.14	2.95
Little Lake	1.86	3.96	15.24	4.20
Cable Lake	1.34	2.77	9.14	3.20
Bond Falls Flowage	8.65	3.05	8.50	NA
Au Train Lake	3.42	2.29	9.14	3.62
Greenwood Reservoir	3.90	2.44	9.14	3.70
Unnamed	17.30	3.35	21.95	NA
Dead River Storage Basin	11.08	3.66	18.00	9.50
Muskallonge Lake	3.19	2.51	6.10	2.90
Rice Lake	2.65	2.59	2.74	1.10
Lac La Belle	4.88	2.74	11.58	6.50
Group 9:Upper Penins	sula, Large, Olig	otrophic		
Bear Lake	1.27	5.53	18.29	8.50
Big Bear Lake	1.39	4.27	10.97	4.70
Thumb Lake	2.06	7.77	46.33	9.50
Lake Esau	1.29	5.03	8.84	4.10
Long Lake	1.54	7.24	18.59	7.10
Lake Antoine	3.00	4.57	7.62	3.60
Unnamed	1.71	4.88	21.34	6.20
Skeels Lake	1.92	4.57	15.20	4.50
Glimmerglass Lake	3.48	7.01	18.29	7.20
Lake Gratiot	5.89	4.72	21.34	8.10
Lake Medora	2.79	4.27	9.14	4.40
Group 10:Northern Lo	ower Peninsula,	Small, Eutrophi	ic	
Brownlee Lake	0.37	1.98	36.44	7.62

Name	Surface Area (km ²)	Secchi Depth (m)	Max Depth (m)	Mean Depth (m)
Group 11:Northern L	ower Peninsula,	Small, Mesotrop	ohic	
Half Moon Lake	0.26	2.82	21.34	6.90
Sand Lake	0.24	2.01	4.57	1.80
Twin Lakes	0.46	2.44	5.79	2.70
Winfield Lake	0.48	2.29	17.37	5.20
Lake Montcalm	0.34	3.52	20.40	9.10
Horseshoe Lake	0.39	2.44	13.41	7.40
Brockway Lake	0.07	2.51	7.93	NA
School Section Lake	0.49	2.90	10.06	3.10
Crystal Lake	0.29	2.13	10.36	4.10
Townline Lake	0.31	3.66	15.85	5.40
Hillsview Lake	0.40	2.29	12.19	3.50
Unnamed	0.47	3.35	6.71	3.40
Bergess Lake	0.24	3.20	14.63	4.60
Merrill Lake	0.34	3.35	8.23	2.60
Pleiness Lake	0.40	3.20	11.60	NA
Idlewild Lake	0.42	2.90	6.71	2.70
Tiff Lake	0.30	2.51	12.19	NA
Five Lake	0.48	2.01	20.42	3.20
Unnamed	0.21	3.66	15.85	NA
Todd Lake	0.33	2.44	15.24	6.50
Windover Lake	0.28	3.66	20.73	4.50
Hackert Lake	0.49	3.51	4.88	2.00
Sunrise Lake	0.31	3.51	20.12	7.50
Unnamed	0.27	3.05	18.29	NA
Harper Lake	0.34	3.96	18.29	5.50
Lake Four	0.24	2.44	12.19	NA
Canfield Lake	0.14	3.02	8.84	NA
Berry Lake	0.28	3.35	8.53	3.10
Goose Lake	0.40	2.74	4.27	1.20
Round Lake	0.36	2.44	4.57	1.50
Long Lake	0.27	2.74	4.57	1.90
Grousehaven Lake	0.35	2.74	16.46	6.70
Devoe Lake	0.48	2.13	16.15	6.20
Vaughn Lake	0.46	3.05	19.81	6.50
Bass Lake	0.38	2.59	5.79	1.90
East Lake	0.38	2.26	6.10	2.00
Herendeene Lake	0.16	3.02	11.28	NA
Cedar Lake	0.21	3.35	18.59	7.50

Name	Surface Area	Secchi Depth	Max Depth	Mean Depth
Ivanic	(km^2)	(m)	(m)	(m)
Wilson Lake	0.36	2.74	14.33	3.50
Ellsworth Lake	0.43	3.05	12.80	5.20
Saint Clair Lake	0.24	3.05	9.75	3.10
Group 12:Northern Lo	wer Peninsula,	Small, Oligotrop	ohic	
Englewright Lake	0.22	5.03	20.12	5.70
Baptist Lake	0.33	4.57	19.81	6.80
Cowden Lake	0.45	4.27	15.24	5.10
Brush Lake	0.08	4.88	8.23	NA
Diamond Lake	0.24	4.27	18.29	5.10
Arnold Lake	0.49	6.10	25.91	NA
Lake George	0.37	4.11	14.63	8.30
North Lake	0.35	4.27	28.04	7.30
Group 13:Northern Lov	wer Peninsula, I	Medium, Eutroj	phic	
Rainbow Lake	0.68	1.98	6.71	1.50
Stony Lake	1.12	1.83	13.11	6.60
Robinson Lake	0.54	1.26	9.14	5.50
Diamond Lake	0.76	1.65	7.62	2.70
Boyles Creek	1.16	1.83	22.30	4.90
Ross Lake	1.02	1.83	5.79	NA
Lincoln Lake	0.70	1.22	4.00	NA
Lincoln Lake	0.70	1.22	4.00	NA
Hicks Lake	0.65	1.07	10.06	3.10
Cranberry Lake	0.66	1.37	6.10	2.40
Hardwood Lake	0.73	1.22	10.67	3.40
Indian Lake	0.87	1.98	4.57	NA
George Lake	0.76	1.68	27.43	6.20
Group 14:Northern Lo	wer Peninsula, I	Medium, Mesot	rophic	
Duck Lake	0.87	3.96	19.80	NA
Muskellunge Lake	0.55	2.44	11.28	5.60
Bills Lake	0.80	3.81	27.43	5.70
Little Whitefish Lake	0.73	3.35	13.11	5.50
Kimball Lake	0.59	2.13	16.15	NA
Pickerel Lake	1.24	2.44	22.25	7.80
McLaren Lake	1.05	3.35	21.34	6.90
Blue Lake	0.90	2.74	15.24	NA
Coldwater Lake	1.16	2.44	19.81	9.90
East Lake	0.83	2.44	15.55	NA
Stevenson Lake	0.55	2.13	14.33	4.10
School Section Lake	0.76	3.05	7.93	3.40

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Name	Surface Area (km ²)	Secchi Depth (m)	Max Depth	Mean Depth
D		. ,	(m)	(m)
Pratt Lake	0.76	3.05	8.53	3.00
Bass Lake	1.19	3.96	13.72	3.10
Long Lake	0.84	3.20	23.17	4.40
Londo Lake	0.73	2.44	4.27	1.80
Sand Lake	0.99	2.44	7.62	2.30
West Londo Lake	0.80	2.90	4.88	2.10
Clear Lake	0.83	3.56	15.24	3.50
Unnamed	0.74	2.29	9.40	2.00
Cedar Lake	0.58	2.44	2.44	1.00
Cedar Hedge Lake	0.78	3.20	20.12	5.80
Brown Bridge Pond	0.69	2.59	8.84	3.40
Jewell Lake	0.75	2.59	10.36	1.80
Cedar Lake	1.00	3.05	14.02	8.00
Lake of the Woods	0.70	2.44	4.27	1.50
Susan Lake	0.51	2.69	4.57	1.80
Group 15:Northern Lo	wer Peninsula,	Medium, Oligot	rophic	
Clear Lake	0.50	5.18	9.14	3.10
Nichols Lake	0.62	4.57	15.24	4.90
Unnamed	0.65	5.18	20.73	9.00
Lake George	0.52	4.11	7.62	2.80
Pine Lake	0.67	5.33	15.24	6.20
Group 16:Northern Lo	wer Peninsula, I	Large, Eutrophi	ic	
White Lake	10.43	1.98	21.30	6.90
White Lake	10.43	1.98	21.30	6.90
Fremont Lake	3.34	1.68	26.82	9.90
Tamarack Lake	1.28	0.76	5.18	1.50
Pentwater Lake	1.98	1.98	12.30	NA
Pentwater Lake	1.98	1.98	12.30	NA
Round Lake	2.19	1.52	3.05	1.40
Hamlin Lake	19.00	1.98	23.99	4.50
Sixmile Lake	1.50	1.83	9.40	4.00
Group 17:Northern Lo				
Big Blue Lake	1.36	3.35	15.24	3.80
Croton Dam Pond	5.02	2.77	12.19	2.70
Hardy Dam Pond	12.75	3.05	33.50	9.40
Rogers Dam Pond	1.36	2.13	9.14	NA
Unnamed	1.91	2.76	12.80	4.00
Unnamed	6.89	2.59	14.30	2.10
	0.07		11.20	NA

Name	Surface Area (km ²)	Secchi Depth (m)	Max Depth (m)	Mean Depth (m)
Pere Marquette Lake	2.47	2.29	11.60	NÁ
Wolf Lake	1.65	3.66	3.96	1.80
Rose Lake	1.51	2.38	9.14	4.70
Manistee Lake	4.23	2.13	14.90	7.50
Lake Missaukee	8.26	2.26	8.23	3.20
Sage Lake	3.15	2.74	24.38	5.80
Hodenpyl Dam Pond	7.52	3.05	21.30	7.00
Little Au Sable Lake	1.29	3.51	13.72	4.00
Long Lake	1.96	3.96	18.90	5.00
Loud Dam Pond	2.41	3.05	9.14	4.30
Foote Dam Pond	6.87	3.81	12.19	5.80
Upper Herring Lake	2.32	2.51	7.93	4.60
Lower Herring Lake	1.81	2.77	18.29	8.90
Lower Herring Lake	1.81	2.77	18.29	8.90
Spider Lake	1.80	3.96	10.97	2.80
Platte Lake	10.25	2.44	27.43	8.23
Lake Ann	2.03	3.27	22.86	9.70
Bass Lake	1.39	3.66	8.80	2.70
Lake Skegemog	11.20	2.74	8.84	3.40
Lime Lake	2.68	3.05	20.42	5.00
Little Traverse Lake	2.61	3.35	16.46	5.30
Clam Lake	1.77	3.05	8.84	4.00
Intermediate Lake	6.36	2.44	6.10	2.60
Unnamed	1.40	3.35	12.50	3.00
Round Lake	1.43	2.13	4.88	1.70
Unnamed	13.92	2.83	18.29	3.05
Group 18:Northern L	ower Peninsula,	Large, Oligotro	phic	
Horsehead Lake	1.79	4.57	16.46	NA
Fife Lake	2.42	4.72	16.76	4.50
Alcona Dam Pond	3.92	4.42	12.19	5.00
Duck Lake	7.87	4.11	29.87	7.30
Silver Lake	2.46	6.40	29.26	6.70
Long Lake	12.01	7.32	26.20	7.90
Boardman Lake	1.29	5.79	22.25	7.90
Birch Lake	1.32	4.27	15.24	6.10
Group 19:Southern Lo	ower Peninsula,	Small, Eutrophi	c	
Round Lake	0.32	1.98	10.67	6.20
Sand Lake	0.39	1.37	6.71	2.50
Morrison Lake	0.47	1.68	7.62	2.60

Name	Surface Area (km ²)	Secchi Depth (m)	Max Depth (m)	Mean Depth (m)
Kenyon Lake	0.30	1.98	8.84	NA
Ackley Lake	0.26	1.07	4.57	1.70
Brandywine Lake	0.30	1.22	7.62	2.90
Newburgh Lake	0.38	0.76	5.50	2.20
Ruppert Lake	0.11	1.76	8.84	NA
Wabascon Lake	0.29	1.22	13.72	6.00
Blind Lake	0.29	1.68	26.52	NA
Hi-Land Lake	0.46	1.37	3.66	NA
Selkirk Lake	0.37	1.83	11.89	3.60
Myers Lake	0.36	1.83	12.50	3.70
Group 20:Southern Lo	ower Peninsula, S	Small, Mesotrop	ohic	
Cub Lake	0.50	3.96	13.11	5.60
Bird Lake	0.47	3.05	19.51	8.30
Carpenter Lake	0.15	3.52	12.19	NA
Bear Lake	0.42	2.44	15.24	4.80
Cary Lake	0.31	3.05	11.58	6.30
South Sand Lake	0.35	2.26	9.75	NA
Belas Lake	0.21	2.59	5.18	1.50
Twin Lakes	0.25	3.51	16.50	5.20
Fish Lake	0.46	3.20	10.97	NA
Deep Lake	0.30	2.90	15.24	7.30
Killarney Lake	0.07	2.13	5.49	NA
Unnamed	0.36	3.35	13.72	5.70
Kelly Lake	0.18	3.96	7.01	NA
Huzzy Lake	0.39	3.20	10.36	4.60
Unnamed	0.35	2.29	7.62	2.40
Paw Paw Lake	0.50	3.51	17.10	7.40
Hogset Lake	0.32	3.35	9.75	NA
Warner Lake	0.23	2.74	9.14	4.10
Unnamed	0.30	2.59	19.50	5.60
Unnamed	0.37	2.90	8.23	3.50
Little Paw Paw Lake	0.41	2.74	8.84	NA
Unnamed	0.16	3.35	10.10	4.60
Rush Lake	0.49	3.96	17.68	5.00
School Section Lake	0.32	3.35	15.24	5.50
South Scott Lake	0.48	3.96	16.76	6.10
North Scott Lake	0.31	2.74	11.28	6.10
Prairie Lake	0.36	2.13	7.32	4.00
Crooked Lake	0.46	3.96	6.10	1.70

Name	Surface Area (km ²)	Secchi Depth (m)	Max Depth (m)	Mean Depth (m)	
Little Cedar Lake	0.30	3.51	8.23	3.10	
Lake Fourteen	0.28	2.13	6.71	2.20	
Green Lake	0.36	2.90	3.35	1.00	
Lake Eleven	0.22	3.35	9.75	5.30	
Clear Lake	0.28	2.13	16.76	7.50	
Appleton Lake	0.23	2.74	11.58	5.60	
Bishop Lake	0.44	3.20	16.46	2.70	
Hall Lake	0.23	3.20	3.66	NA	
Lower Pettibone Lake	0.42	3.05	12.50	5.80	
Carter Lake	0.24	2.74	7.62	2.60	
Leach Lake	0.44	2.59	12.19	6.50	
Wildwood Lake	0.48	2.59	5.18	2.50	
Crooked Lake	0.27	2.74	20.73	NA	
Big Fish Lake	0.43	2.74	21.34	5.00	
Davison Lake	0.23	3.96	20.42	7.60	
Lake Minnawanna	0.24	3.20	6.10	2.30	
Woodard Lake	0.28	3.66	6.71	2.80	
Crockery Lake	0.42	3.35	16.46	7.50	
Baldwin Lake	0.25	3.02	10.70	1.80	
Half Moon Lake	0.21	3.20	14.63	NA	
Unnamed	0.22	3.02	18.30	7.00	
Group 21:Southern Lo	wer Peninsula, S	Small, Oligotrop	ohic		
Lake Lavine	0.36	5.79	21.64	6.50	
Harwood Lake	0.48	4.88	16.76	7.20	
North Sand Lake	0.24	4.78	12.19	NA	
Upper Jeptha Lake	0.24	4.02	21.34	NA	
Crooked Lake	0.33	4.11	12.19	NA	
Long Lake	0.31	4.27	14.94	7.90	
Valley Lake	0.15	4.27	13.72	2.48	
Group 22:Southern Lo	wer Peninsula, I	Medium, Eutroj	phic		
Shavehead Lake	1.21	1.52	21.34	NA	
Long Lake	0.86	1.51	12.80	3.50	
South Lake	0.63	1.37	5.49	1.30	
Round Lake	0.83	0.61	8.23	2.30	
Lake of the Woods	1.21	1.98	10.36	4.50	
Vandercook Lake	0.59	1.68	12.80	7.00	
Maple Lake	0.67	1.22	4.57	2.20	
Fourmile Lake	1.04	1.68	5.49	1.40	
Duck Lake	0.56	1.07	11.89	4.50	

Name	Surface Area (km ²)	Secchi Depth (m)	Max Depth (m)	Mean Depth (m)	
Base Line Lake	0.86	1.52	13.40	9.50	
Joslin Lake	0.89	1.98	6.10	1.60	
Bristol Lake	0.57	1.52	12.19	6.20	
Middle Straits Lake	0.74	1.83	15.24	3.70	
Big Lake	0.88	1.07	4.27	1.20	
Duncan Lake	0.53	1.37	16.76	7.60	
Park Lake	0.74	1.76	8.23	NA	
Kearsley Reservoir	0.74	0.46	6.20	1.50	
Group 23:Southern L	ower Peninsula,	Medium, Mesot	rophic		
Baldwins Lake	1.06	3.12	16.76	NA	
Gilead Lake	0.56	2.13	14.94	5.90	
Omena Lake	0.53	2.90	17.07	5.50	
Paradise Lake	0.75	2.90	17.07	7.30	
Stone Lake	0.64	2.29	17.07	6.06	
Donnell Lake	0.99	2.90	19.20	7.60	
Hemlock Lake	0.61	2.01	19.81	NA	
Clear Lake	0.94	2.59	9.45	3.60	
Pleasant Lake	1.04	2.74	16.15	5.90	
Dewey Lake	0.90	2.29	15.24	NA	
Oliverda Lake	0.62	2.29	10.67	NA	
Gravel Lake	1.20	2.90	15.55	5.60	
Cedar Lake	1.11	3.35	25.60	7.30	
Unnamed	0.61	2.29	12.19	2.40	
Sugarloaf Lake	0.60	3.05	9.14	3.30	
Gourdneck Lake	0.89	2.74	15.85	NA	
Lee Lake	0.53	3.81	14.33	7.90	
Eagle Lake	0.79	3.51	3.05	1.52	
Unnamed	0.76	3.35	12.50	5.80	
Van Auken Lake	1.02	2.44	14.02	6.70	
Sherman Lake	0.60	3.35	10.97	4.10	
Sugarloaf Lake	0.72	2.29	6.10	1.00	
Mill Lake	0.53	3.35	7.62	1.70	
Saddle Lake	1.14	3.05	9.75	2.50	
North Lake	0.91	2.44	17.68	3.93	
Osterhout Lake	0.70	2.59	9.14	1.70	
Eagle Lake	0.88	3.35	17.98	6.50	
South Lake	0.82	3.05	25.30	5.60	
Bruin Lake	0.53	2.44	14.60	5.60	
Base Line Lake	1.02	3.51	19.51	9.50	

Name	Surface Area (km ²)	Secchi Depth (m)	Max Depth (m)	Mean Depth (m)
Clear Lake	0.97	3.35	4.88	2.00
Big Lake	0.61	3.05	9.14	3.90
Long Lake	1.06	2.44	14.63	5.50
West Crooked Lake	0.74	2.29	5.79	NA
East Crooked Lake	1.01	2.44	12.19	4.00
Woodland Lake	1.05	2.90	10.67	2.10
Wolverine Lake	1.16	2.90	17.37	2.30
Thompson Lake	1.07	2.13	15.85	3.00
Middle Lake	0.55	2.29	9.14	5.90
Tipsico Lake	1.00	2.77	8.23	3.00
Green Lake	1.20	2.74	21.03	9.20
Deer Lake	0.56	2.44	19.20	NA
Seven Lakes	0.68	2.51	16.15	NA
Big Pine Island Lake	0.79	2.59	13.72	4.60
Wolf Lake	0.91	2.44	11.60	NA
Bass Lake	0.77	3.35	6.10	2.30
Clifford Lake	0.79	2.74	13.72	4.70
Dickerson Lake	0.92	2.59	14.63	5.90
Murphy Lake	0.74	3.05	12.50	2.00
Group 24:Southern Lo	wer Peninsula,	Medium, Oligot	rophic	
Thompson Lake	0.60	4.88	9.14	4.60
Fish Lake	0.62	4.57	22.86	7.40
Threemile Lake	1.05	4.42	10.67	3.10
Fish Lake	0.61	4.57	17.07	9.00
Cedar Island Lake	0.68	4.11	21.95	8.40
Camp Lake	0.55	4.27	15.20	7.50
Group 25:Southern Lo	wer Peninsula,	Large, Eutrophi	ic	
Unnamed	1.92	0.74	7.32	4.00
Matteson Lake	1.27	1.98	11.58	4.60
North Lake	2.06	1.07	10.67	5.40
Union Lake	2.20	0.91	4.88	2.20
Barton Lake	1.36	1.37	16.20	6.10
Austin Lake	4.46	1.52	3.35	1.30
Belleville Lake	5.05	0.61	9.14	6.10
Unnamed	2.03	1.98	3.66	NA
Unnamed	6.88	0.61	6.10	3.30
Kent Lake	4.21	1.51	10.70	1.90
Thornapple Lake	1.69	1.37	9.45	5.30
Stony Creek Lake	2.02	0.76	7.01	2.40

Name	Surface Area (km ²)	Secchi Depth (m)	Max Depth (m)	Mean Depth (m)
Lake Ponemah	1.63	1.83	24.38	NÁ
Morrison Lake	1.35	0.76	10.97	3.90
Lake Ovid	1.50	0.75	3.05	2.30
C S Mott Lake	2.43	0.76	5.49	1.70
Unnamed	2.63	0.46	12.80	4.00
Unnamed	18.70	1.52	12.50	4.55
Group 26:Southern L	ower Peninsula,	Large, Mesotro	phic	
Klinger Lake	3.36	2.44	21.95	6.40
Lake of the Woods	1.57	3.35	23.17	5.90
Diamond Lake	4.21	2.59	19.51	4.60
Corey Lake	2.42	3.81	24.40	8.19
Palmer Lake	2.95	2.51	11.28	4.15
Round Lake	2.06	3.35	20.42	NA
Fishers Lake	1.34	2.29	12.80	4.00
Devils Lake	5.31	2.29	19.20	4.30
Fish Lake	1.35	2.74	14.33	6.20
Sand Lake	2.21	3.81	16.15	3.40
Magician Lake	2.05	2.59	15.20	3.87
Vineyard Lake	2.19	2.59	12.80	4.20
Bankson Lake	1.47	3.51	12.80	4.10
Clark Lake	2.33	3.81	15.24	3.40
Indian Lake	3.20	3.51	22.90	8.50
Paw Paw Lake	3.73	3.51	27.43	8.80
Long Lake	2.03	2.74	17.37	2.60
Center Lake	3.43	3.81	8.53	0.90
Unnamed	1.35	3.05	9.14	1.40
Grass Lake	1.43	3.20	4.00	0.80
Goguac Lake	1.44	3.35	20.12	5.30
Unnamed	2.42	2.74	15.24	4.00
Gull Lake	8.29	2.13	33.53	NA
Whitmore Lake	2.33	3.96	21.03	4.20
Fine Lake	1.31	2.44	13.72	2.90
Zukey Lake	1.74	3.02	14.63	NA
Pine Lake	2.48	3.20	10.36	3.60
Crooked Lake	2.72	2.26	14.60	2.20
Unnamed	1.46	3.20	15.24	NA
Hutchins Lake	1.53	2.44	10.36	3.20
Miner Lake	1.33	3.51	25.30	7.20
Lake Chemung	1.26	3.96	21.34	8.70

Name	Surface Area (km ²)	Secchi Depth (m)	Max Depth (m)	Mean Depth (m)
Cass Lake	5.31	3.35	37.49	9.10
White Lake	2.33	2.74	9.75	3.30
Pontiac Lake	2.56	2.79	10.36	1.20
Unnamed	1.85	2.74	19.51	3.40
Lobdell Lake	3.18	3.96	23.77	2.20
Lake Orion	2.00	2.90	24.40	5.00
Wabasis Lake	1.64	2.74	17.40	7.60
Group 27:Southern	Lower Peninsula,	Large, Oligotro	phic	
Orchard Lake	3.50	4.27	33.53	6.70
Union Lake	1.92	4.57	33.53	8.70
Lake Fenton	3.56	4.11	27.43	6.20
Lakeville Lake	1.80	5.03	20.10	3.00

B. Validation Results

Table B.1 is a complete summary of the regression-based equivalence testing results using observed meteorological data. For each lake, the minimum region of similarity for the slope and intercept, and the percent of bootstrap iterations where the slope and intercept fell within the region of similarity.

			SLOPE				INTERCEPT				
Lake	- Parameter	Confi Inte		Regio Simila		Region of Similarity Min.		dence erval		on of larity	Region of Similarity Min.
	-	Lower	Upper	Lower	Upper	%	Lower	Upper	Lower	Upper	°C
Alloguash Laka	Ts	0.92	1.03	0.85	1.15	-0.07	14.59	15.56	13.70	15.70	0.87
Allequash Lake	Tm	0.98	1.15	0.85	1.15	0.70	13.67	14.81	13.12	15.12	0.17
Big	Ts	0.90	1.00	0.85	1.15	-0.13	13.72	15.10	13.62	15.62	-0.90
Muskellunge Lake	Tm	0.78	0.95	0.85	1.15	-0.21	11.41	12.61	12.25	14.25	-1.88
Converted Labo	Ts	0.89	1.00	0.85	1.15	-0.13	13.48	14.84	13.63	15.63	-1.17
Crystal Lake	Tm	0.99	1.14	0.85	1.15	0.17	11.44	12.32	11.01	13.01	-0.62
Created Dec	Ts	0.86	1.04	0.85	1.15	-0.13	14.23	15.54	14.05	16.05	-0.87
Crystal Bog	Tm	0.98	1.14	0.85	1.15	0.13	14.61	15.54	13.17	15.17	1.40
Suculation of Tala	Ts	0.90	1.00	0.85	1.15	-0.11	13.81	14.95	13.49	15.49	-0.73
Sparkling Lake	Tm	1.12	1.25	0.85	1.15	0.27	11.54	12.23	10.24	12.24	0.97
Trout Lake	Ts	0.81	0.96	0.85	1.15	-0.21	12.87	14.48	12.83	14.83	-1.04
TIOUL Lake	Tm	0.82	0.97	0.85	1.15	-0.22	9.85	10.87	9.18	11.18	0.71
Fish Lake	Ts	0.93	1.09	0.85	1.15	0.10	12.99	14.74	14.78	16.78	-2.72
FISH Lake	Tm	0.98	1.19	0.85	1.15	0.19	9.84	10.89	10.79	12.79	-1.96
Laka Wingro	Ts	1.00	1.21	0.85	1.15	0.19	14.93	16.33	16.28	18.28	-2.43
Lake Wingra	Tm	0.59	0.82	0.85	1.15	-0.39	11.14	12.57	15.99	17.99	-5.89
Lake Monona	Ts	0.95	1.12	0.85	1.15	0.12	14.58	16.05	15.64	17.65	-2.19
Lake Monona	Tm	0.73	0.86	0.85	1.15	-0.27	10.89	11.78	14.04	16.04	-4.19
Lake Mendota	Ts	0.99	1.09	0.85	1.15	0.11	13.58	14.82	13.81	15.81	-1.25
Lake Menuola	Tm	0.79	0.89	0.85	1.15	-0.21	10.23	10.95	11.71	13.71	-2.52

Table B.1. Summary of regression-based equivalence test for each validation lake using observed climate data.

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Table B.2 is a complete summary of the regression-based equivalence testing results using modeled meteorological data. For each lake, the minimum region of similarity for the slope and intercept, and the percent of bootstrap iterations where the slope and intercept fell within the region of similarity.

		SLOPE					INTERCEPT				
Lake	Parameter	Confidence Interval		Region of Similarity		Region of Similarity Min.	Confidence Interval		Region of Similarity	Region of Similarity Min.	
		Lower	Upper	Lower	Upper	%	Lower	Upper	Lower	Upper	°C
Allequash Lake	Ts	0.92	1.12	0.80	1.20	0.15	16.33	18.18	10.70	18.70	3.52
	Tm	0.96	1.21	0.80	1.20	0.23	15.19	16.85	10.12	18.12	2.74
Big Muskellunge Lake	e Ts	0.94	1.10	0.80	1.20	0.15	16.17	17.99	10.62	18.62	3.46
	Tm	0.70	0.86	0.80	1.20	-0.28	11.33	12.26	9.25	17.25	-1.99
Crystal Lake	Ts	0.97	1.13	0.80	1.20	0.18	16.15	17.94	10.63	18.63	3.43
	Tm	0.89	1.07	0.80	1.20	0.09	11.33	12.18	8.01	16.01	-0.76
Crystal Bog	Ts	0.70	1.07	0.80	1.20	-0.26	16.27	18.86	11.05	19.05	3.87
	Tm	0.85	1.20	0.80	1.20	0.20	16.43	18.64	10.17	18.17	4.56
Sparkling Lake	Ts	0.96	1.16	0.80	1.20	0.21	16.18	18.23	10.49	18.49	3.78
	Tm	0.97	1.16	0.80	1.20	0.17	11.37	12.17	7.24	15.24	0.93
Trout Lake	Ts	0.91	1.09	0.80	1.20	0.15	15.63	17.56	9.83	17.83	3.75
	Tm	0.87	1.01	0.80	1.20	-0.15	10.56	11.41	6.18	14.18	1.27
Fish Lake	Ts	0.96	1.16	0.80	1.20	0.17	18.18	19.96	11.78	19.78	4.16
	Tm	0.79	1.04	0.80	1.20	-0.19	11.74	12.83	7.79	15.79	1.10
Lake Wingra	Ts	0.87	1.22	0.80	1.20	0.20	19.19	21.72	13.28	21.28	4.49
	Tm	0.63	1.04	0.80	1.20	-0.35	16.97	19.33	13.00	21.00	2.51
Lake Monona	Ts	0.92	1.13	0.80	1.20	0.14	19.07	20.82	12.65	20.65	4.20
	Tm	0.51	0.72	0.80	1.20	-0.47	12.22	13.45	11.04	19.04	-2.83
Lake Mendota	Ts	0.96	1.13	0.80	1.20	0.14	18.27	19.86	10.81	18.81	5.08
	Tm	0.62	0.77	0.80	1.20	-0.37	12.00	12.87	8.71	16.71	-0.70

Table B.2. Summary of regression-based equivalence test for each validation lake using modeled climate data.

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C. Supplementary Materials

File Name: 01_Validation_Climate_Data

Description: Daily observed and modeled meteorological data for the northern and southern validation site.

File Name: 02_Validation_Lake_Data Description: Observed lake temperature data for all validation lakes.

File Name: 03 Future Climate Data

Description: Daily modeled meteorological data for 2020 to 2099 for each of the climate regions.

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Using Google Maps, Google Earth and Street View: https://www.google.com/permissions/geoguidelines.html

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Below, you'll find information on:

General guidelines

Uses in print

Uses on the web and in applications

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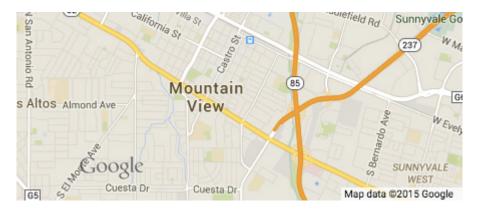
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